

## SCHNEIDERHÖHNITE, $\text{Fe}^{2+}\text{Fe}^{3+}\text{As}_5^{3+}\text{O}_{13}$ , A DENSELY PACKED ARSENITE STRUCTURE

FRANK C. HAWTHORNE

Department of Earth Sciences, University of Manitoba, Winnipeg, Manitoba R3T 2N2

### ABSTRACT

The crystal structure of schneiderhöhnite,  $\text{Fe}^{2+}\text{Fe}^{3+}\text{As}_5^{3+}\text{O}_{13}$ ,  $a = 8.924(2)$ ,  $b = 10.016(3)$ ,  $c = 9.103(3) \text{ \AA}$ ,  $\alpha = 59.91(2)$ ,  $\beta = 112.41(2)$ ,  $\gamma = 81.69(2)^\circ$ ,  $V = 590.8(3) \text{ \AA}^3$ , space group P1, has been solved by direct methods and refined to a residual  $R$  of 4.5% on 3184 unique observed [ $I > 2.5\sigma(I)$ ] reflections of graphite-monochromated MoK $\alpha$  radiation. Five unique As positions, each with trigonal pyramidal co-ordination, form two ( $\text{As}_2\text{O}_5$ ) dimers and one ( $\text{AsO}_3$ ) monomer. In five unique Fe positions, each with distorted pseudo-octahedral co-ordination, the observed  $\langle \text{Fe}-\text{O} \rangle$  distances show four of the octahedra to be occupied by  $\text{Fe}^{3+}$  and the fifth by  $\text{Fe}^{2+}$ . The octahedra polymerize by sharing edges and corners, which unite with the arsenite pyramids to form a complex, densely packed framework structure. A brief survey of recently refined arsenite minerals shows a fairly constant geometry for the arsenite group, with  $\langle \text{As}-\text{O} \rangle = 1.776(12) \text{ \AA}$  and  $\langle \text{O}-\text{As}-\text{O} \rangle = 97.2(2.0)^\circ$ .

**Keywords:** crystal structure, arsenite, schneiderhöhnite, Tsumeb.

### SOMMAIRE

La structure cristalline de la schneiderhöhnite,  $\text{Fe}^{2+}\text{Fe}^{3+}\text{As}_5^{3+}\text{O}_{13}$ ,  $a = 8.924(2)$ ,  $b = 10.016(3)$ ,  $c = 9.103(3) \text{ \AA}$ ,  $\alpha = 59.91(2)$ ,  $\beta = 112.41(2)$ ,  $\gamma = 81.69(2)^\circ$ ,  $V = 590.8(3) \text{ \AA}^3$ , groupe spatial P1, a été résolue par méthodes directes et affinée jusqu'au résidu  $R = 4.5\%$  sur 3184 réflexions observées uniques [ $I > 2.5\sigma(I)$ ] en radiation MoK $\alpha$  monochromatisée sur graphite. Cinq positions uniques d'As, chacune à coordination pyramidale trigonale, forment deux dimères  $\text{As}_2\text{O}_5$  et un monomère  $\text{AsO}_3$ . Dans cinq positions uniques de Fe, chacune avec un octaèdre difforme comme polyèdre de coordination, les distances  $\langle \text{Fe}-\text{O} \rangle$  observées indiquent que quatre des pseudo-octaèdres sont occupés par  $\text{Fe}^{3+}$  et le cinquième, par  $\text{Fe}^{2+}$ . Les octaèdres se polymérisent en mettant en commun arêtes et sommets, lesquels s'unissent aux pyramides trigonales pour former une structure complexe en charpente à empilement compact. Une brève revue des arsenites minéraux dont la structure a été affinée récemment montre que le groupe des arsenites est caractérisé par une géométrie assez constante, avec  $\langle \text{As}-\text{O} \rangle = 1.776(12) \text{ \AA}$  et  $\langle \text{O}-\text{As}-\text{O} \rangle = 97.2(2.0)^\circ$ .

(Traduit par la Rédaction)

**Mots-clés:** structure cristalline, arsenite, schneiderhöhnite, Tsumeb.

### INTRODUCTION

Schneiderhöhnite is an iron arsenite first reported

by Ottemann *et al.* (1973). It occurs at Tsumeb, S.W. Africa, in a druse in the deep oxidation zone, associated with chalcocite and zincian stottite.

### EXPERIMENTAL

The crystal used in this work is from the Tsumeb copper deposit, S.W. Africa, and was obtained from the Department of Mineralogy and Geology, Royal Ontario Museum, collection number M33238. The crystal was mounted on a Nicolet R3m automated four-circle diffractometer, and twenty-five intense reflections were centred using graphite-monochromated MoK $\alpha$  X radiation. Least-squares refinement of the setting angles produced the cell dimensions given in Table 1, together with an orientation matrix relating the crystal axes to the diffractometer axes. A total of 3833 reflections was measured over one asymmetric unit out to a maximum  $2\theta$  of  $60^\circ$  according to the experimental procedure of Hawthorne (1985). An empirical absorption-correction ( $\psi$ -scan) was applied, together with corrections for Lorentz, polarization and background effects; of the 3469 unique reflections, 3187 were classed as observed [ $I > 2.5\sigma(I)$ ].

### STRUCTURE SOLUTION AND REFINEMENT

Scattering curves for neutral atoms together with anomalous dispersion corrections were taken from Cromer & Mann (1968) and Cromer & Liberman (1970), respectively.  $R$  indices are of the form given in Table 1 and are expressed as percentages. The SHELXTL system of programs was used for the computational procedures.

The structure was solved by direct methods. The phase set with the maximum combined figure of merit gave a satisfactory solution, and all atomic positions were derived from the initial model and a difference-Fourier map calculated after the first few cycles of refinement. For an isotropic thermal model, refinement converged to an  $R$  index of 6.4%. Conversion to anisotropic temperature-factors and refinement of all variables converged to an  $R$  index of 4.9% (observed) and an  $R_w$  index of 6.1% (observed). Final positional parameters are given in Table 2, anisotropic temperature-factor coefficients in Table 3, selected interatomic distances and angles in Tables 4 and 5, and a bond-valence analysis is offered in Table 6. Observed and calculated

structure-factors are available from the Depository of Unpublished Data, CISTI, National Research Council of Canada, Ottawa, Ontario K1A 0S2.

## DISCUSSION

### Description of the structure

There are five unique As sites with trigonal pyramidal co-ordination, this co-ordination resulting

from the  $4s^2$  stereoactive lone pair of electrons on each  $\text{As}^{3+}$  cation. The variation in both individual and mean bond-lengths in the  $(\text{As}^{3+}\text{O}_3)$  groups lies within the normal range exhibited by natural and synthetic arsenites. There are five unique Fe sites, all with pseudo-octahedral co-ordination. The variation in mean octahedral bond-lengths is compatible with  $\text{Fe}(1) = \text{Fe}(2) = \text{Fe}(3) = \text{Fe}(4) = \text{Fe}^{3+}$  and  $\text{Fe}(5) = \text{Fe}^{2+}$ , as indicated by sum of the relevant ionic radii:  $\text{Fe}(1) = \text{Fe}(2) = \text{Fe}(3) = \text{Fe}^{3+} - \text{III}^{\text{O}} =$

TABLE 1. MISCELLANEOUS INFORMATION: SCHNEIDERHÖHNITE

a	8.924(2) $\text{\AA}$	Crystal size (mm)	0.08x0.28x0.30
b	10.016(3)	Rad/Mono	Mo/Gr
c	9.103(3)	Total  Fo	3673
$\alpha$	59.91(2) $^\circ$	Fo  > 4 $\sigma$ .	3184
B	112.41(2)	Final R(obs)	4.5%
$\gamma$	81.69(2)	Final R <sub>w</sub> (obs)	4.9%
V	590.8(3) $\text{\AA}^3$	$\mu(\text{cm}^{-1})$	197
Space Group	P $\bar{1}$		
Unit cell contents:	$2[\text{Fe}^{2+}\text{Fe}^{3+}_3\text{As}^{3+}_5\text{O}_{13}]$		
R	$= \Sigma( \text{Fo}  -  \text{Fc} )/\Sigma \text{Fo} $		
R <sub>w</sub>	$= [\sum_w ( \text{Fo}  -  \text{Fc} )^2 / \sum_w \text{Fo}^2]^{1/2}$ , w = 1		

TABLE 2. ATOMIC PARAMETERS FOR SCHNEIDERHÖHNITE

	x	y	z	$U_{\text{equiv.}}^*$
As(1)	0.8020(1)	0.3006(1)	0.9938(1)	0.48(3)
As(2)	0.6696(1)	0.1689(1)	0.4947(1)	0.45(3)
As(3)	0.8843(1)	0.9023(1)	0.1653(1)	0.41(3)
As(4)	0.7322(1)	0.7385(1)	0.8569(1)	0.43(3)
As(5)	0.7555(1)	0.5380(1)	0.5125(1)	0.62(3)
Fe(1)	0	0	1/2	0.42(6)
Fe(2)	1/2	1/2	0	0.43(6)
Fe(3)	0.8861(1)	0.3823(1)	0.3186(1)	0.33(5)
Fe(4)	0.4546(1)	0.1967(1)	-0.0798(1)	0.48(5)
Fe(5)	0.6363(1)	0.8945(1)	0.3863(1)	0.55(5)
O(1)	0.7959(6)	0.5808(6)	0.3174(7)	0.7(2)
O(2)	0.4326(6)	0.2002(6)	0.3831(7)	0.7(1)
O(3)	0.3994(6)	0.1247(6)	0.1315(7)	0.8(2)
O(4)	0.8763(7)	0.1068(6)	0.0244(7)	0.8(2)
O(5)	0.7526(6)	0.0298(6)	0.4512(7)	0.7(3)
O(6)	0.9205(6)	0.2214(6)	0.2470(7)	0.6(2)
O(7)	0.5559(6)	0.7309(6)	0.3323(7)	0.7(2)
O(8)	0.0046(6)	0.8515(6)	0.4090(7)	0.7(2)
O(9)	0.7131(6)	0.5508(6)	0.0133(7)	0.6(2)
O(10)	0.1213(6)	0.3717(6)	0.3916(7)	0.8(3)
O(11)	0.6498(6)	0.3551(6)	0.2939(7)	0.7(2)
O(12)	0.6656(6)	0.9617(6)	0.1256(7)	0.5(2)
O(13)	0.4059(6)	0.6856(6)	0.0210(7)	0.6(1)

\*  $U_{\text{equiv.}} = U_{\text{equiv.}} \times 10^2$

TABLE 3. ANISOTROPIC TEMPERATURE-FACTOR COEFFICIENTS FOR SCHNEIDERHÖHNITE

	$U_{11}^*$	$U_{22}$	$U_{33}$	$U_{23}$	$U_{13}$	$U_{12}$
As(1)	34(3)	48(3)	45(3)	-17(3)	25(3)	-12(2)
As(2)	19(3)	53(3)	44(3)	-19(3)	19(3)	-12(2)
As(3)	11(3)	51(3)	55(3)	-31(3)	17(2)	-5(2)
As(4)	9(3)	50(3)	53(3)	-19(3)	20(2)	-11(2)
As(5)	50(3)	53(3)	61(3)	-26(3)	24(3)	-22(2)
Fe(1)	8(6)	36(5)	54(6)	-18(5)	11(5)	2(5)
Fe(2)	7(6)	42(5)	47(6)	-16(5)	6(5)	-4(5)
Fe(3)	1(4)	29(4)	36(4)	-10(3)	-0(3)	4(3)
Fe(4)	26(4)	53(4)	50(4)	-27(4)	15(4)	-14(3)
Fe(5)	49(4)	58(4)	67(4)	-39(4)	34(4)	55(5)
O(1)	68(2)	77(21)	60(22)	-42(19)	35(19)	-18(18)
O(2) <sup>†</sup>	1(15)	83(15)	41(13)	0(0)	0(0)	0(0)
O(3)	71(23)	63(21)	86(23)	-36(19)	49(20)	9(18)
O(4)	81(23)	43(20)	60(22)	-13(19)	40(19)	1(18)
O(5)	40(11)	73(21)	141(25)	-73(20)	62(20)	-38(18)
O(6)	45(21)	57(20)	50(21)	-28(18)	22(18)	-17(17)
O(7)	34(21)	81(21)	78(22)	-40(19)	34(19)	-33(18)
O(8)	44(22)	56(20)	37(21)	-16(18)	-20(18)	-4(17)
O(9)	24(21)	64(20)	75(23)	-31(19)	26(19)	-20(17)
O(10)	26(21)	112(22)	107(24)	-70(21)	25(19)	-33(18)
O(11)	41(21)	42(20)	67(22)	-6(18)	16(19)	-12(17)
O(12)	13(20)	66(20)	68(22)	-38(19)	20(18)	-17(17)
O(13) <sup>‡</sup>	4(15)	37(13)	57(13)	0(0)	0(0)	0(0)

<sup>†</sup> $U_{1,j} = U_{1,j} \times 10^4$ ; <sup>‡</sup>non-positive definite

TABLE 4. SELECTED BOND-LENGTHS ( $\text{\AA}$ ) IN SCHNEIDERHÖHNITE

As(1)-0(4)a	1.797(7)	As(4)-0(2)e	1.805(5)	
As(1)-0(6)a	1.768(6)	As(4)-0(3)e	1.744(7)	
As(1)-0(13)e	1.788(6)	As(4)-0(9)a	1.784(5)	
<As(1)-0>	<u>1.784</u>	<As(4)-0>	<u>1.778</u>	
As(2)-0(2)	1.858(5)	As(5)-0(1)	1.799(7)	
As(2)-0(5)	1.767(8)	As(5)-0(7)	1.787(4)	
As(2)-0(11)	1.758(5)	As(5)-0(10)e	1.771(7)	
<As(2)-0>	<u>1.794</u>	<As(5)-0>	<u>1.786</u>	
As(3)-0(4)c	1.761(5)			
As(3)-0(8)d	1.796(6)			
As(3)-0(12)	1.790(6)			
<As(3)-0>	<u>1.782</u>			
Fe(1)-0(5)f,g	2.022(6)	x2	Fe(4)-0(1)b	2.164(4)
Fe(1)-0(6)f,g	1.976(4)	x2	Fe(4)-0(3)	1.974(7)
Fe(1)-0(8)h,i	2.036(7)	x2	Fe(4)-0(7)b	1.991(7)
<Fe(1)-0>	<u>2.011</u>		Fe(4)-0(12)h	2.050(4)
Fe(2)-0(9),b	2.014(6)	x2	Fe(4)-0(12)b	2.157(6)
Fe(2)-0(11),b	1.967(5)	x2	Fe(4)-0(13)b	1.943(6)
Fe(2)-0(13),b	2.055(7)	x2	<Fe(4)-0>	<u>2.047</u>
<Fe(2)-0>	<u>2.012</u>		Fe(5)-0(2)c	2.200(7)
Fe(3)-0(1)	2.027(7)		Fe(5)-0(3)c	2.158(4)
Fe(3)-0(6)	2.044(7)		Fe(5)-0(5)c	2.001(7)
Fe(3)-0(8)e	2.064(4)		Fe(5)-0(7)	2.079(7)
Fe(3)-0(9)	2.056(4)		Fe(5)-0(10)e	2.338(4)
Fe(3)-0(10)d	1.913(6)		Fe(5)-0(12)	2.253(7)
Fe(3)-0(11)	2.106(6)		<Fe(5)-0>	<u>2.172</u>
<Fe(3)-0>	<u>2.035</u>			



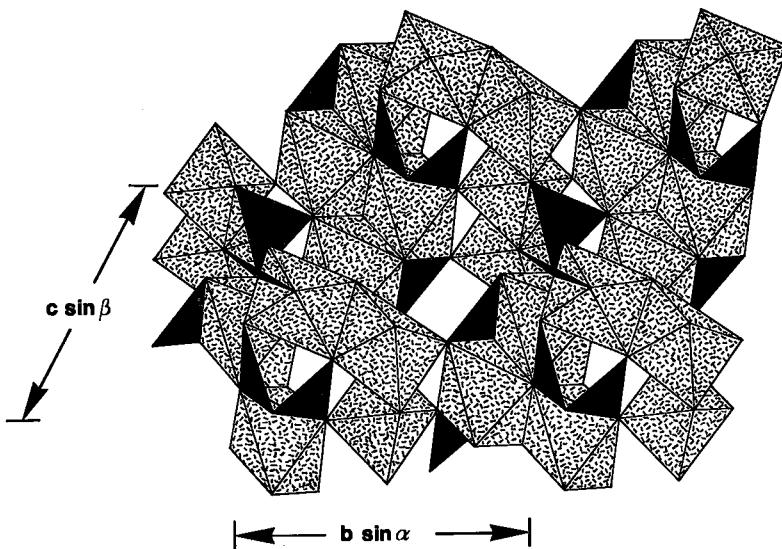


FIG. 1. The crystal structure of schneiderhöhnite, viewed down the X axis; Fe octahedra are hatched,  $\text{AsO}_3$  groups are shown in black.

TABLE 7. BOND LENGTHS ( $\text{\AA}$ ) AND ANGLES ( $^\circ$ ), MEAN OXYGEN COORDINATION NUMBER (C.N.), BOND LENGTH ( $\Delta$ ) AND ANGLE ( $\sigma^2$ ) DISTORTION PARAMETERS, AND MEAN CATION ELECTRONEGATIVITY ( $\langle x \rangle$ ) FOR ARSENITE GROUPS FROM PRECISELY REFINED MINERAL STRUCTURE

	As-O	As-O	As-O	$\langle \text{As}-\text{O} \rangle$	O-As-O	O-As-O	$\langle \text{O}-\text{As}-\text{O} \rangle$	C.N.	* $\Delta$	* $\sigma^2$	$\langle x \rangle$	Ref.
Reinerite	1.753(5)	1.772(3)	1.772(3)	1.766	97.6(1)	97.6(1)	100.0(1)	98.4	3.00	0.26	1.28	1.86 (1)
"	1.772(4)	1.771(3)	1.771(3)	1.771	95.2(1)	95.2(1)	100.0(1)	96.8	3.00	0.00	5.12	1.86 "
Cafarsite	1.78(2)	1.78(2)	1.78(2)	1.78	96.4(8)	96.4(8)	96.4	4.00	0.00	0.00	1.75	(2)
"	1.74(1)	1.74(1)	1.79(1)	1.76	97.4(5)	93.8(5)	96.0(5)	95.7	4.33	1.83	2.20	1.75 "
"	1.73(1)	1.75(2)	1.80(1)	1.76	92.3(7)	101.2(5)	95.5(8)	96.3	3.67	2.80	13.55	1.75
Stenuggarite	1.778(7)	1.775(7)	1.776(6)	1.776	99.6	99.0	91.3	96.6	3.30	0.01	14.28	1.85
"	1.736(7)	1.826(7)	1.801(6)	1.788	92.7	101.3	94.3	96.1	3.33	4.50	13.95	1.85
Claudetite	1.79(1)	1.81(1)	1.72(1)	1.77	98.9(6)	103.3(6)	93.7(6)	98.6	2.00	4.74	15.40	2.18
"	1.77(1)	1.79(1)	1.75(1)	1.77	94.9(6)	98.0(6)	89.7(6)	94.2	2.00	0.85	11.73	2.18 "
Asbecasite	1.787(7)	1.757(8)	1.846(7)	1.797	102.8(3)	94.9(3)	90.1(3)	95.9	3.33	4.23	27.42	1.62
Dixenite	1.779(5)	1.779(5)	1.779(5)	1.779	94.8(2)	94.8(2)	94.8(2)	94.8	3.00	0.00	0.00	1.76
"	1.755(5)	1.755(5)	1.755(5)	1.755	98.0(2)	98.0(2)	98.0(2)	98.0	4.00	0.00	0.00	1.76 "
"	1.721(5)	1.762(5)	1.779(5)	1.754	98.1(3)	103.8(2)	106.6(3)	102.8	3.67	1.93	12.51	1.76
Armangite	1.760(3)	1.782(3)	1.795(3)	1.779	91.7(2)	92.5(2)	103.3(2)	95.8	4.00	0.66	27.98	1.82
"	1.745(3)	1.772(3)	1.814(4)	1.777	92.3(2)	100.0(2)	106.0(2)	99.4	3.67	2.55	31.44	1.82 "
"	1.773(3)	1.776(4)	1.809(4)	1.786	93.6(2)	94.8(2)	100.6(2)	96.3	3.67	0.83	9.34	1.82 "
Paulmooreite	1.747(9)	1.750(9)	1.826(9)	1.774	95.7(4)	98.3(5)	99.9(4)	98.0	3.67	4.25	3.00	2.26
"	1.733(9)	1.772(8)	1.842(9)	1.782	95.4(4)	100.6(5)	98.7(4)	98.3	3.00	6.40	4.62	2.26 "
Hematolite	1.79(1)	1.79(1)	1.79(1)	1.79	94.2(5)	94.2(5)	94.2(5)	94.2	4.00	0.00	0.00	1.66
Magnussonite	1.750(6)	1.758(6)	1.779(6)	1.762	94.0(3)	94.0(3)	107.0(3)	98.3	4.00	0.48	37.56	1.70 (10)
Schneiderhöhnite	1.797(7)	1.768(6)	1.788(6)	1.784	99.8(2)	101.2(3)	94.6(3)	98.5	2.67	0.46	8.06	2.02
"	1.852(5)	1.767(8)	1.758(5)	1.794	99.2(3)	99.6(2)	96.9(3)	98.6	3.00	5.58	1.42	2.02 "
"	1.761(5)	1.796(8)	1.790(5)	1.782	98.2(3)	98.8(2)	97.3(3)	98.1	3.00	0.74	0.11	2.02 "
"	1.805(5)	1.744(6)	1.784(5)	1.778	98.1(3)	98.5(3)	97.7(3)	98.1	3.00	2.03	0.11	2.02 "
"	1.799(7)	1.787(4)	1.771(7)	1.786	89.3(3)	101.7(3)	93.4(3)	94.8	3.00	0.41	26.61	2.02 "

References: (1) Ghose et al. (1977); (2) Edenharder et al. (1977); (3) Coda et al. (1977); (4) Pertlik (1978); (5) Cannillo et al. (1969); (6) Araki & Moore (1981); (7) Moore & Araki (1979a); (8) Araki et al. (1980); (9) Moore & Araki (1978); (10) Moore & Araki (1979b); (11) This study.

$$\ast \Delta = \frac{\sum ((l_i - \langle l_1 \rangle)^2 / \langle l_1 \rangle^2)}{3}; \quad \sigma^2 = \frac{\sum ((O-\text{As}-O) - \langle O-\text{As}-O \rangle)^2}{3}$$

tion; this often leads to a lower precision on the bond lengths. In addition, inadequacies in the absorption correction commonly give rise to inaccurate bond-lengths. Consequently, it is possible that some of the observed variations in Table 7 arise from experimental error, and obscure any structural effects that may affect the geometry of the arsenite group.

#### ACKNOWLEDGEMENTS

It is a pleasure to acknowledge the co-operation of Toronto Joe Mandarino and Fojo Wicks, curators of the mineral collection at the Royal Ontario Museum, for supplying the crystal of schneiderhöhnite used in this study. This work was supported by

a fellowship, an operating grant and a major equipment grant from the Natural Sciences and Engineering Research Council of Canada.

## REFERENCES

- ARAKI, T. & MOORE, P.B. (1981): Dixenite,  $\text{Cu}^{1+}\text{Mn}^{2+}_{14}\text{Fe}^{3+}(\text{OH})_6(\text{As}^{3+}\text{O}_3)_5(\text{Si}^{4+}\text{O}_4)_2(\text{As}^{5+}\text{O})_4$ : metallic  $[\text{As}^{3+}_4\text{Cu}^{1+}]$  clusters in an oxide matrix. *Amer. Mineral.* **66**, 1263-1273.
- \_\_\_\_\_, \_\_\_\_ & BRUNTON, G.D. (1980): The crystal structure of paulmooreite,  $\text{Pb}_2[\text{As}_2\text{O}_5]$ : dimeric arsenite groups. *Amer. Mineral.* **65**, 340-345.
- BAUR, W.H. (1974): The geometry of polyhedral distortions. Predictive relationships for the phosphate group. *Acta Cryst.* **B30**, 1195-1215.
- \_\_\_\_\_(1978): Variation of mean Si-O bond lengths in silicon-oxygen tetrahedra. *Acta Cryst.* **B34**, 1751-1756.
- BROWN, I.D. (1981): The bond-valence method: an empirical approach to chemical structure and bonding. In *Structure and Bonding in Crystals*, II (M. O'Keeffe & A. Navrotsky, eds.). Academic Press, New York.
- CANNILLO, E., GIUSEPPETTI, G. & TADINI, C. (1969): The crystal structure of asbecasite. *Acad. Naz. Lincei, Rend. Sci. fis. mat. nat. Ser. VIII, XLVI*, 457-467.
- CODA, A., DAL NEGRO, A., SABELLI, C. & TAZZOLI, V. (1977): The crystal structure of stenhuggarite. *Acta Cryst.* **B33**, 1807-1811.
- CROMER, D.T. & LIBERMAN, D. (1970): Relativistic calculation of anomalous scattering factors for X-rays. *J. Chem. Phys.* **53**, 1891-1898.
- \_\_\_\_\_, & MANN, J.B. (1968): X-ray scattering factors computed from numerical Hartree-Fock wave functions. *Acta Cryst.* **A24**, 321-324.
- EDENHARTER, A., NOWACKI, W. & WEIBEL, M. (1977): Zur Struktur und Zusammensetzung von Cafarsit, ein As(III)-oxid, kein Arsenat. *Schweiz. mineral. petrog. Mitt.* **57**, 1-16.
- GHOSE, S., BOVING, P., LACHAPELLE, W.A. & WAN, C. (1977): Reinerite,  $\text{Zn}_3(\text{AsO}_3)_2$ : an arsenite with a novel type of Zn-tetrahedral double chain. *Amer. Mineral.* **62**, 1129-1134.
- HAWTHORNE, F.C. (1985): Refinement of the crystal structure of bloedite: structural similarities in the  $[\text{V}^{\text{IV}}\text{M}(\text{T}\Phi_4)\Phi_n]$  finite cluster minerals. *Can. Mineral.* **23**, 669-674.
- \_\_\_\_\_, & FAGGIANI, R. (1979): Refinement of the structure of descloizite. *Acta Cryst.* **B35**, 717-720.
- MOORE, P.B. & ARAKI, T. (1978): Hematolite: a complex dense-packed sheet structure. *Amer. Mineral.* **63**, 150-159.
- \_\_\_\_\_, & \_\_\_\_ (1979a): Armangite,  $\text{Mn}^{2+}_{26}[\text{As}^{3+}_6(\text{OH})_4\text{O}_1][\text{As}^{3+}_6\text{O}_{18}]_2[\text{CO}_3]$ , a fluorite derivative structure. *Amer. Mineral.* **64**, 748-757.
- \_\_\_\_\_, & \_\_\_\_ (1979b): Magnussonite, manganese arsenite, a fluorite derivative structure. *Amer. Mineral.* **64**, 390-401.
- OTTEMANN, J., NUBER, B. & GEIER, B.H. (1973): Schneiderhöhnit ein natürliches Eisen-Arsen-Oxid aus der tiefen Oxidationszone von Tsumeb. *Neues Jahrb. Mineral. Monatsh.*, 517-523.
- PERTLIK, F. (1978): Verfeinerung der Kristallstruktur des Minerals Claudetit,  $\text{As}_2\text{O}_3$  ("Claudetit I"). *Monatsh. Chem.* **109**, 277-282.
- SHANNON, R.D. (1975): Systematic studies of interatomic distances in oxides. In *The Physics and Chemistry of Minerals and Rocks* (R.G.J. Strens, ed.). J. Wiley & Sons, London.

*Received December 3, 1984, revised manuscript accepted April 17, 1985.*