# Magnesio-hornblende from Lüderitz, Namibia: mineral description and crystal chemistry

ROBERTA OBERTI<sup>1,\*</sup>, MASSIMO BOIOCCHI<sup>2</sup>, FRANK C. HAWTHORNE<sup>3</sup> AND MARCO E. CIRIOTTI<sup>4</sup>

- <sup>1</sup> CNR-Istituto di Geoscienze e Georisorse, Sede secondaria di Pavia, via Ferrata 1, I-27100 Pavia, Italy
- <sup>2</sup> Centro Grandi Strumenti, Università di Pavia, via Bassi 21, I-27100 Pavia, Italy
- Department of Geological Sciences, University of Manitoba, Winnipeg, Manitoba, R3T 2N2, Canada
- <sup>4</sup> Associazione Micromineralogica Italiana, via San Pietro 55, I-10073 Devesi-Cirié, Italy

[Received 16 October 2017; Accepted 14 November 2017; Associate Editor: Anthony Kampf]

# **ABSTRACT**

**KEYWORDS:** magnesio-hornblende, electron-microprobe analysis, optical properties, powder-diffraction pattern, crystal-structure refinement, Lüderitz, Namibia.

#### Introduction

THE generic name 'hornblende' is used widely by petrologists and mineral collectors for any black undefined Al-bearing amphibole in the calcium (and sometimes in the sodium–calcium) subgroups. Therefore, the currently approved scheme for amphibole nomenclature and classification (Hawthorne *et al.*, 2012) uses the rootname hornblende for the charge arrangement <sup>A</sup>□<sup>B</sup>Ca<sub>2</sub> <sup>C</sup>(R<sub>4</sub><sup>2</sup>+R<sup>3+</sup>)<sup>T</sup>(Si<sub>7</sub>Al)O<sub>22</sub> <sup>W</sup>(OH,F,Cl)<sub>2</sub> where Al is dominant over R<sup>3+</sup> cations and the prefix 'magnesio-' indicates Mg dominance over R<sup>2+</sup> cations (exceptions are allowed only for riebeckite, arfvedsonite, hastingsite and hornblende).

Despite its widespread occurrence (~220 compositions listed in the mindat.org web page: https://www.mindat.org/min-2524.html, some of which should not be classified as magnesio-hornblende), the type description of an amphibole with composition corresponding to  ${}^{A}\Box {}^{B}Ca_{2}{}^{C}(Mg_{4}AI)^{T}(Si_{7}AI)$   $O_{22}{}^{W}(OH)_{2}$  has never been published.

In the most recent list of minerals approved by the International Mineralogical Association (IMA version July 2017; https://www.ima-mineralogy.org/Minlist.htm), magnesio-hornblende is designated as a redefined valid species (by Hawthorne et al., 2012) with unknown type locality and without an original reference. Therefore, it is a potential new species which requires formal approval by the IMA Commission on New Minerals, Nomenclature and Classification (CNMNC) and publication of the mineral data. This work describes magnesio-hornblende from Lüderitz, Namibia approved by IMA-CNMNC vote 2017-059.

\*E-mail: oberti@crystal.unipv.it https://doi.org/10.1180/minmag.2017.081.093

# Occurrence and optical properties

The specimen under investigation was found in 1970 by MEC in the sand dunes of the Lüderitz, Karas Region, Namibia (~26°38′52″S, 15°09′28″E). Lüderitz is a surreal colonial relic, a true 19th century Bayarian village at the southern margin of the Namib Desert coast. It was founded in 1883 as a trading post, and in 1909 enjoyed a sudden but short period of prosperity when diamonds were discovered nearby. The Namib Desert was part of the passive margin during Early Cretaceous rifting of the South Atlantic, and extends for nearly 600 km along the Namibian coast between Lüderitz ( $\sim$ 27°S) and the Kuiseb River ( $\sim$ 23°S); it is 100 to 150 km wide, extending inland from the coast to the base of the Great Escarpment at the 1000 m contour. The sand sea (~34,000 km<sup>2</sup>) is dominated by large linear dunes, with areas of star-shaped dunes on its eastern margin and a belt of simple and compound transverse and barchanoid dunes (Fig. 1) along the coast (Garzanti et al., 2012). Dune sands in the coastal Namib are invariably lithofeldspatho-quartzose volcaniclastic, and have a homogeneous composition from Lüderitz to Walvis Bay. They were originally deposited by the Orange River, which has its source 3000 km away in the basaltic rift highlands of Lesotho and flows into the ocean 300 km south of Lüderitz, to which sands are dragged by vigorous longshore currents (Garzanti *et al.*, 2012). Plagioclase exceeds K-feldspar. In the sand, mafic volcanic rock fragments predominate over granitoid, sedimentary (quartzose to feldspatho-quartzose siltstone/sandstone, shale, minor limestone and hybrid carbonate) and metamorphic (quartz-sericite, quartz-mica, quartz-epidote, amphibolite) grains. Samples enriched in heavy minerals are usually dominated by clinopyroxene, with subordinate



Fig. 1. Barchanoid dunes near Lüderitz.

opaque Fe-Ti-Cr oxides, garnet, epidote and amphibole. Landward of Lüderitz, the quartz content increases at the expense of volcanic detritus; garnet, Fe-Ti-Cr oxides, amphibole (including green and green-brown hornblende), epidote, titanite, staurolite and zircon also become more significant, suggesting contributions from the Koichab River and the locally exposed Namaqua basement rocks (Garzanti et al., 2012).

The holotype specimen (Fig. 2) consists of a friable block of subhedral to anhedral magnesiohornblende crystals up to a few mm in size. The crystals are green to dark green with good cleavage parallel to {110}. A spindle stage was used to orient a crystal for measurement of refractive indices and 2V by extinction curves (Bartelmehs et al., 1992). The optical orientation was determined by transferring the crystal from the spindle stage to a singlecrystal diffractometer and measuring the relative axial relations by X-ray diffraction. In transmitted plane-polarized light, magnesio-hornblende is pleochroic (X= pale yellow, Y= bluish green, Z= dark green) with Z > Y > X. Magnesio-hornblende is non fluorescent and has a vitreous lustre. It is biaxial (-), with  $\alpha = 1.640(2)$ ,  $\beta = 1.654(2)$ ,  $\gamma =$ 1.666(2) (measured with gel-filtered Na light,  $\lambda =$ 589.9 nm), 2V (meas.) =  $82(1)^{\circ}$  and 2V (calc.) = 84.9°. The dispersion is weak (v > r), and the orientation is:  $X^{\wedge} a = 33.7^{\circ}$  (in  $\beta$  obtuse),  $Y \parallel b, Z^{\wedge}$  $c = 18.2^{\circ}$  (in  $\beta$  acute). The density calculated based on the chemical formula and the unit-cell parameters is 3.137 g·cm<sup>-3</sup>. The compatibility index  $[1 - (K_p/K_c)]$  (Mandarino, 2007) is 0.010 (superior).

The holotype material is deposited in the mineralogical collections of the Museo di Mineralogia, Sistema Museale di Ateneo,



Fig. 2. The holotype specimen studied in this work. Scale in cm.

## MAGNESIO-HORNBLENDE

TABLE 1. Atom coordinates, refined site-scattering values\* (ss, epfu), and equivalent atom-displacement parameters  $(B_{eq}, Å^2)$  for magnesio-hornblende crystal 1325.

Site	SS	x/a	y/b	z/c	$B_{ m eq}$
O(1)		0.1090(2)	0.08794(11)	0.2154(4)	0.76(4)
O(2)		0.1194(2)	0.17303(10)	0.7295(4)	0.70(4)
O(3)	16.04(12)	0.1108(3)	0	0.7149(5)	0.81(7)
O(4)	` ´	0.3680(2)	0.24863(11)	0.7904(4)	0.94(4)
O(5)		0.3482(2)	0.13658(11)	0.1022(4)	0.95(4)
O(6)		0.3435(2)	0.11837(11)	0.5944(4)	0.93(4)
O(7)		0.3361(3)	0	0.2858(6)	1.06(6)
T(1)		0.28034(7)	0.08481(4)	0.29774(14)	0.45(2)
T(2)		0.29007(7)	0.17183(4)	0.80727(13)	0.51(2)
M(1)	31.56(14)	0	0.08892(5)	1/2	0.61(2)
M(2)	29.73(14)	0	0.17739(5)	0	0.52(2)
M(3)	16.40(6)	0	0	0	0.57(3)
M(4)	37.8(2)	0	0.27851(5)	1/2	0.77(2)
M(4')	2.34(10)	0	0.2549(10)	1/2	1.0
A	0.89(8)	0	1/2	0	2.0
A(m)	0.84(11)	0.028(6)	1/2	0.081(14)	2.0
A(2)	1.67(11)	0	0.478(2)	0	2.0
H	1.96	0.196(6)	0	0.758(10)	1.0

<sup>\*</sup>Hawthorne et al. (1995); 'epfu' - electrons per formula unit.

TABLE 2. Selected interatomic distances (Å), angles (°), tetrahedral and octahedral angle variances (TAV, OAV, °^2) and quadratic elongations (TQE, OQE) according to Robinson *et al.* (1971) refined for magnesio-hornblende crystal 1325.

T(1)-O(1) T(1)-O(5) T(1)-O(6) T(1)-O(7) <t(1)-o> TAV TQE</t(1)-o>	1.630(2) 1.657(2) 1.651(2) 1.6337(12) 1.643 4.40 1.0010	T(2)-O(2) T(2)-O(4) T(2)-O(5) T(2)-O(6) <t(2)-o> TAV TQE</t(2)-o>	1.623(2) 1.598(2) 1.648(2) 1.668(2) 1.634 19.50 1.0049	M(4)–O(2) ×2 M(4)–O(4) ×2 M(4)–O(5) ×2 M(4)–O(6) ×2 < M(4)–O>	2.400(2) 2.304(2) 2.723(2) 2.545(2) 2.493
M(1)-O(1) ×2 M(1)-O(2) ×2 M(1)-O(3) ×2 < $M(1)$ -O> OAV OQE	2.060(2) 2.106(2) 2.105(2) 2.090 40.78 1.0126	M(2)-O(1) ×2 M(2)-O(2) ×2 M(2)-O(4) ×2 < $M(2)$ -O> OAV OQE	2.107(2) 2.072(2) 1.992(2) 2.057 22.51 1.0072	M(3)-O(1) ×4 M(3)-O(3) ×2 <m(3)-o> OAV OQE</m(3)-o>	2.086(2) 2.075(3) 2.082 58.09 1.0180
A-O(5) ×4 A-O(6) ×4 A-O(7) ×2 <a-o> M(4')-O(2) ×2 M(4')-O(4) ×2 M(4')-O(5) ×2 M(4')-O(6) ×2 <m(4')-o></m(4')-o></a-o>	3.003(2) 3.142(2) 2.477(3) 2.953 2.077(13) 2.253(2) 2.985(12) 2.872(14) 2.547	A(m)-O(5) ×2 A(m)-O(5) ×2 A(m)-O(6) ×2 A(m)-O(7) A(m)-O(7) A(m)-O(7) A(m)-O(9) M(1)- $M(2)$	3.05(3) 3.02(3) 2.83(5) 2.40(5) 3.28(7) 2.63(6) 2.90 3.213(2)	A(2)-O(5) ×2 A(2)-O(6) ×2 A(2)-O(7) ×2 < A(2)-O> T(1)-O(5)- $T(2)T(1)$ -O(6)- $T(2)T(1)$ -O(7)- $T(1)O(5)-O(6)-O(5)O(6)-O(7)-O(6)$	2.69(3) 2.89(2) 2.507(6) 2.70 135.48(12) 138.52(12) 139.4(2) 165.71(12) 105.81(13)

Table 3. Powder X-ray diffraction data (*d* in Å) for magnesio-hornblende crystal 1325.

$I_{\rm rel}$	d(calc)	h k l	$I_{\mathrm{rel}}$	d(calc)	h k l
22.9	9.033	020	39.2	2.164	2 6 1
74.2	8.412	110	5.4	2.137	Ī 52
8.7	5.088	1 3 0	20.0	2.047	2 0 2
22.3	4.892	$\bar{1} \ 1 \ 1$	27.5	2.017	3 5 1
16.2	4.516	0 4 0			$\bar{4} \ 0 \ 2$
5.3	4.206	220	8.4	2.002	3 7 0
16.4	3.884	$\bar{1}$ 3 1	5.2	1.965	1 9 0
48.5	3.386	131	6.3	1.890	5 1 0
29.1	3.274	2 4 0	5.8	1.877	$\bar{4}$ 61
72.5	3.121	3 1 0	11.7	1.867	Ī 91
33.9	2.941	2 2 1	5.6	1.848	Ī 72
11.5	2.804	3 3 0	5.0	1.843	$\bar{4}$ 4 2
31.0	2.731	$\bar{3} \ 3 \ 1$	5.0	1.812	5 3 0
100.0	2.709	151	5.9	1.747	5 12
45.4	2.596	061	13.6	1.689	$\bar{1}$ 3 3
57.5	2.541	$\bar{2} 0 2$			$\bar{2} \ 8 \ 2$
8.4	2.382	3 5 0	8.2	1.678	0 2 3
40.7	2.338	$\bar{3}  5  1$	26.7	1.650	4 6 1
20.3	2.322	$\bar{4} \ 2 \ 1$	8.1	1.637	4 8 0
14.8	2.301	$\bar{1} 7 1$	12.4	1.618	1110
20.7	2.280	$\bar{3} 1 2$	28.1	1.583	6 0 0
7.8	2.215	$\bar{2} 4 2$			Ī 53
7.9	2.182	171	6.9	1.556	4 0 2

The strongest eight reflections are in bold. Only peaks with  $I_{\rm rel} \ge 5$  are reported.

University of Pavia, under the catalogue number 2017-01. The refined and analysed crystal of this work has the code 1325 in the amphibole database of the CNR-IGG Pavia.

# Single-crystal diffraction and powderdiffraction analysis

The crystal studied is  $0.60~\mathrm{mm} \times 0.45~\mathrm{mm} \times 0.10~\mathrm{mm}$  in size. The unit-cell parameters are a=9.8308(7), b=18.0659(11), c=5.2968(4) Å,  $\beta=104.771(6)^\circ$  and V=909.64 (11) ų, and were calculated by a least-squares procedure using the  $d^*$  values measured for reflections belonging to  $60~\mathrm{selected}$  rows of reciprocal space and occurring in the  $\theta$  range  $-30^\circ$  to  $+30^\circ$ . Two monoclinic equivalents were collected in the  $2\theta$  range  $4-60^\circ$  by a Philips PW1100 4-circle diffractometer using MoK $\alpha$  radiation; absorption and Lorentz and polarization corrections were applied. Merging of equivalent reflections gave  $R_{\mathrm{sym}}=2.60\%$  on  $1290~\mathrm{pairs}$ . Reflections with  $I>3\sigma_I$  were considered as observed during unweighted full-

matrix least-squares refinement on F using a program locally written to deal with complex solid-solutions (Cannillo et al., 1983). Scattering curves for fully ionized chemical species were used at sites where chemical substitutions occur; neutral vs. ionized scattering curves were used at the T and at the anion sites [except O(3)]. More details on the refinement strategy can be found in Oberti et al. (1992) and Oberti et al. (2017). The final disagreement indexes are  $R_{\text{obs}} = 2.72\%$  (1244 reflections) and  $R_{\text{all}} = 3.15\%$ (1380 reflections). Refined atom coordinates and displacement parameters, and selected bond lengths and angles are given in Tables 1 and 2, respectively. Crystallographic information files with embedded structure factors have been deposited with the Principal Editors of Mineralogical Magazine and are available as supplementary material (see below). The a:b:c ratio calculated from the unit-cell parameters is 0.544:1:0.293.

Powder X-ray diffraction data ( $CuK\alpha$ ,  $\lambda$ = 1.54178 Å) were obtained using the *XPREP* utility of *SAINT* (Bruker, 2003), which generates a 2D powder diffractogram (Debye-Scherrer technique) starting from the  $F_{\rm obs}^2$  collected on the single crystal and taking into account solely the information concerning the unit-cell dimensions and the Laue symmetry. No Lorentz and polarization corrections were applied. Data are given in Table 3.

## Chemical analysis

Chemical analysis (10 points) on the crystal used for structure refinement was undertaken with a Cameca SX-100 electron microprobe (WDS mode, 15 kV, 20 nA, count time 20 s and 5 µm beam diameter). The standards used are as follows: Si and Ca: diopside (TAP); Ti: titanite (LPET); Al: andalusite (TAP); Cr: chromite (LPET); Fe: fayalite (LLiF); Mn: spessartine (LLiF); Mg: forsterite (LTAP); Zn: gahnite (LLiF); Ni: pentlandite (LLiF); Na: albite (TAP); K: orthoclase (LPET); F: fluoro-riebeckite (TAP) and Cl: tugtupite (LPET). H<sub>2</sub>O was estimated based on 2 = (OH,F,Cl) atoms per formula unit (apfu) and taking into account the constraints on the group sites based on stoichiometry and sitescattering values (for A, B and C cations) obtained from the structure refinement; this method allows calculation of Fe3+/Fe2+. The oxide wt.% and the calculated unit-formula are reported in Table 4. The proposed empirical formula for crystal 1325 is:  $\begin{array}{l} \text{From Na}_{0.22}\text{K}_{0.05})_{\Sigma 1.00}^{\text{B}}(\textbf{Ca}_{\textbf{1.79}}\text{Fe}_{0.10}^{2+}\text{Mg}_{0.04}\text{Mn}_{0.03}^{2+}\\ \text{Na}_{0.04})_{\Sigma 2.00}^{\text{C}}(\textbf{Mg}_{\textbf{3.48}}\text{Fe}_{0.97}^{2+}\text{Al}_{0.28}\text{Fe}_{0.23}^{3+}\text{Cr}_{0.01}^{3+}\\ \text{Ti}_{0.03})_{\Sigma 5.00}^{\text{T}}(\text{Si}_{7.18}\text{Al}_{0.82})_{\Sigma 8.00}\text{O}_{22}^{\text{W}}(\textbf{OH}_{\textbf{1.93}}\text{F}_{0.05}) \end{array}$ 

#### MAGNESIO-HORNBLENDE

Table 4. Chemical composition (10 analytical points), unit formula (based on 24 anions and 15.27 cations) and a comparison between observed and calculated site scattering for magnesio-hornblende (1325).

Oxide	Wt.%(esd)	Range	Apfu			
SiO <sub>2</sub>	50.24(23)	49.16-51.10	Si	7.18	Ca	1.79
TiO <sub>2</sub>	0.24(2)	0.22 - 0.25	Al	0.82	Na	0.04
$Al_2\tilde{O}_3$	6.52(7)	6.26 - 6.72	Sum T	8.00	$Fe^{2+}$	0.10
$Fe_2O_3$	2.17				$Mn^{2+}$	0.03
$Cr_2O_3$	0.10(2)	0.08 - 0.11	Ti <sup>4+</sup>	0.03	Mg	0.04
$V_2O_3$	0.03(4)	0.00 - 0.05	$Al^{3+}$	0.28	Sum B	2.00
FeO	8.87		Fe <sup>3+</sup>	0.23		
[FeO] <sub>tot</sub>	[10.82(20)]	10.69-10.98	$V^{3+}$	_	Na	0.22
MgO	16.52(9)	16.44-16.59	Cr <sup>3+</sup>	0.01	K	0.05
MnO	0.25(5)	0.23 - 0.28	Mg	3.48	Sum A	0.27
NiO	0.03(5)	0.00 - 0.06	$Mn^{2+}$	_		
ZnO	0.02(6)	0.00 - 0.07	Fe <sup>2+</sup>	0.97	$(OH)^{-}$	1.93
CaO	11.68(12)	11.64-11.75	Ni	_	$F^{-}$	0.05
Na <sub>2</sub> O	0.92(6)	0.87 - 1.01	Zn	_	C1	0.02
$K_2O$	0.30(2)	0.23 - 0.35	Sum C	5.00	Sum W	2.00
$H_2O^*$	2.02			Group site sc	attering (epfu)	
F	0.11(30)	0.00 - 0.31		obs (SREF)		calc (EMP)
Cl	0.10		C	77.69		77.50
O = F	-0.05		В	40.14		40.07
O = C1	-0.02		A	3.40		3.37
Total	100.05		Total	121.23		120.94

<sup>\*</sup> calculated based on 24 (O, OH, F, Cl) with (OH + F + Cl) = 2 apfu.

Table 5. Site populations for magnesio-hornblende, crystal 1325. There is close agreement between the refined values of site-scattering (ss, epfu) and mean bond lengths (mbl, Å) and those calculated based on the proposed site-populations\*.

Site	Site population (apfu)		ss (epfu)		mbl (Å)	
		Refined	Calc.	Refined	Calc.	
<i>T</i> (1)	3.18 Si + 0.82 Al			1.643	1.643	
T(2)	4.00 Si			1.634		
M(1)	$1.46 \text{ Mg} + 0.54 \text{ Fe}^{2+}$	31.56	31.56	2.090	2.090	
M(2)	$1.37 \text{ Mg} + 0.14 \text{ Fe}^{2+} + 0.22 \text{ Al} + 0.23 \text{ Fe}^{3+} + 0.01 \text{ Cr}^{3+} + 0.03 \text{ Ti}^{4+}$	29.73	29.82	2.057	2.056	
M(3)	$0.65 \text{ Mg} + 0.29 \text{ Fe}^{2+} + 0.06 \text{ Al}$	16.40	16.12	2.082	2.082	
ΣC cations		77.69	77.50			
B cations	$1.79 \text{ Ca} + 0.10 \text{ Fe}^{2+} + 0.03 \text{ Mn}^{2+} + 0.04 \text{ Mg} + 0.04 \text{ Na}$	40.14	40.07			
A cations	0.22 Na + 0.05 K	3.40	3.37			
W anions	1.93 (OH) <sup>-</sup> + 0.05 F <sup>-</sup> + 0.02 Cl <sup>-</sup>	16.04	16.23			

<sup>\*</sup>Hawthorne et al. (1995)

 $Cl_{0.02}]_{\Sigma 2.00}$  (where the dominant cations/anions relevant to nomenclature issues are in bold). The simplified end-member formula of

 $\begin{array}{ll} \text{magnesio-hornblende is } \ ^{A} \square ^{B} \text{Ca}_{2} \ ^{C} (\text{Mg}_{4} \text{Al})^{T} (\text{Si}_{7} \text{Al}) \\ \text{O}_{22} \ ^{W} (\text{OH})_{2} \ \text{which requires SiO}_{2} \ 51.68, \ \text{Al}_{2} \text{O}_{3} \ 12.53, \\ \text{MgO } 19.80, \ \text{CaO } 13.78, \ \text{H}_{2} \text{O } 2.21, \ \text{total } 100.00 \ \text{wt.}\%. \end{array}$ 

#### ROBERTA OBERTI ETAL.

TABLE 6. A comparison of the optical and crystallographic properties available for magnesio-hornblende (this work), magnesio-ferri-fluoro-hornblende (Oberti et al., 2016a), and ferro-ferri-hornblende (Oberti et al., 2016b).

	Magnesio-hornblende <sup>1</sup>	Magnesio-ferri-fluoro- hornblende <sup>2</sup>	Ferro-ferri-hornblende <sup>3</sup>
Colour	Dark green	Dark brown	Dark green
Optical class	Biaxial (–)	Biaxial (–)	Biaxial (-)
Pleochroism	X = pale yellow	X= pale grey	X = medium gold/brown
	Y = bluish green	Y = dark grey	Y = dark brown/black
	Z = dark green	Z = pale brownish grey	Z = dark grey
Orientation	$X \wedge a = 33.7^{\circ}$ ( $\beta$ obtuse)	$X \wedge a = 47.6^{\circ}$ ( $\beta$ obtuse)	$X \wedge a = 26.2^{\circ} (\beta \text{ obtuse})$
	$Y \parallel b$	$Y \parallel b$	$Y \parallel b$
	$Z \wedge c = 18.2^{\circ}$ ( $\beta$ acute)	$Z \wedge c = 33.4^{\circ}$ ( $\beta$ acute)	$Z \wedge c = 11.5^{\circ}$ ( $\beta$ acute)
α	1.640(2)	1.669(2)	1.697(2)
β	1.654(2)	1.676(2)	1.722(5)
γ	1.666(2)	1.678(2)	1.726(5)
2V <sub>meas</sub> (°)	85(1)°	74(1)°	35.7(4)°
Space group	C2/m	C2/m	C2/m
a (Å)	9.8308(7)	9.839(5)	9.9307(5)
b (Å)	18.0659(11)	18.078(9)	18.2232(10)
c (Å)	5.2968(4)	5.319(3)	5.3190(3)
β (°)	104.771(6)	104.99(3)	104.857(1)
$V(\mathring{A}^3)$	909.64(11)	913.9(9)	930.40(9)
<t(1)-o> (Å)</t(1)-o>	1.643	1.642	1.642
< T(2) - O > (A)	1.634	1.636	1.635
<m(1)-o> (Å)</m(1)-o>	2.090	2.078	2.117
< M(2) - O > (A)	2.057	2.067	2.086
< M(3) - O > (A)	2.082	2.076	2.111
< M(4) - O > (Å)	2.493	2.502	2.514
T(1)–O(5)– $T(2)$ (°)	135.5	136.4	136.7
T(1)-O(6)- $T(2)$ (°)	138.5	139.8	139.1
T(1)-O(7)- $T(1)$ (°)	139.4	140.0	141.3
O(5)-O(6)-O(5)(°)	165.7	167.6	167.8
O(6)-O(7)-O(6) (°)	105.8	108.5	108.3

 $<sup>{}^{1} \</sup>stackrel{A}{(\square_{0.73} Na_{0.22} K_{0.05})} \stackrel{B}{\Sigma_{1.00}} (Ca_{1.79} Fe_{0.10}^{2+} Mn_{0.03}^{2+} Mg_{0.04} Na_{0.04}) \stackrel{C}{\Sigma_{0.00}} (Mg_{3.48} Fe_{0.97}^{2+} Al_{0.28} Fe_{0.23}^{3+} Cr_{0.01}^{3+} Ti_{0.03}) \stackrel{C}{\Sigma_{5.00}} (Ga_{3.48} Fe_{0.97}^{2+} Al_{0.28} Fe_{0.23}^{3+} Cr_{0.01}^{3+} Ti_{0.03}) \stackrel{C}{\Sigma_{5.00}} (Ga_{3.48} Fe_{0.97}^{3+} Al_{0.28} Fe_{0.23}^{3+} Cr_{0.01}^{3+} Ti_{0.03}) \stackrel{C}{\Sigma_{5.00}} (Ga_{3.48} Fe_{0.23}^{3+} Al_{0.28} Fe_{0.23}^{3+} Cr_{0.01}^{3+} Ti_{0.03}) \stackrel{C}{\Sigma_{5.00}} (Ga_{3.48} Fe_{0.23}^{3+} Al_{0.28} Fe_{0.23}^{3+} Cr_{0.01}^{3+} Ti_{0.03}) \stackrel{C}{\Sigma_{5.00}} (Ga_{3.48} Fe_{0.23}^{3+} Al_{0.28} Fe_{0.23}^{3+} Cr_{0.28}^{3+} Ti_{0.28} Fe_{0.23}^{3+} Cr_{0.28}^{3+} Ti_{0.28}^{3+} Fe_{0.23}^{3+} Ti_{0.28}^{3+} Ti_{0.28}^{3+} Fe_{0.23}^{3+} Ti_{0.28}^{3+} Fe_{0.23}^{3+} Ti_{0.28}^{3+} Ti_{0.2$ 

# The crystal chemistry of magnesio-hornblende

Site populations were obtained taking into account the experimental mean bond lengths and sitescattering values for the individual sites (Hawthorne et al., 1995; Oberti et al., 2007). The results are reported in Table 5. As expected, <sup>T</sup>Al is ordered at the T(1) site, smaller B cations (Mn,Fe,Mg) are ordered at the M(4') site and  ${}^{C}R^{3+}$  cations are almost ordered at the M(2) site but for a small amount of Al occurring at M(3). This was assumed to improve the agreement between the observed and calculated mean bond lengths at the M(2) and M(3)sites, is in accord with the high Mg content of the sample and may suggest high-T crystallization (Oberti et al., 1995).

A comparison of the optical and crystallographic properties of the three members related to the hornblende rootname that have been characterized recently is provided in Table 6. In these three hornblende compositions the 2V value drops

 $<sup>\</sup>begin{array}{l} T(\underbrace{Si_{0.89}Al_{1.11}}_{0.77}Na_{0.10}K_{0.13})_{\Sigma 1.00} \underbrace{V(F_{1.56}(OH)_{0.37}Cl_{0.08})_{\Sigma 2.00}}_{E(Ca_{1.93}Na_{0.07})_{\Sigma 2.00}}(Mg_{1.16}Fe_{3.21}^{2+}Mn_{0.06}Fe_{0.45}^{3+}Al_{0.12}Ti_{0.01})_{\Sigma 5.01} \\ T(\underbrace{Si_{7.26}Al_{0.74})_{\Sigma 8.00}O_{22}}_{E(OH)_{0.39}F_{0.01}Cl_{0.10})_{\Sigma 2.00}}(Mg_{1.16}Fe_{3.21}^{2+}Mn_{0.06}Fe_{0.45}^{3+}Al_{0.12}Ti_{0.01})_{\Sigma 5.01} \\ T(\underbrace{Si_{7.26}Al_{0.74}}_{E(OH)_{0.39}F_{0.01}Cl_{0.10})_{\Sigma 2.00}}(Mg_{1.16}Fe_{3.21}^{2+}Mn_{0.06}Fe_{0.45}^{3+}Al_{0.12}Ti_{0.01})_{\Sigma 5.01} \\ T(\underbrace{Si_{7.26}Al_{0.74}}_{E(OH)_{0.39}F_{0.01}Cl_{0.10})_{\Sigma 2.00}}(Mg_{1.16}Fe_{3.21}^{2+}Mn_{0.06}Fe_{0.45}^{3+}Al_{0.12}Ti_{0.01})_{\Sigma 5.01} \\ T(\underbrace{Si_{7.26}Al_{0.74}}_{E(OH)_{0.39}F_{0.01}Cl_{0.10})_{\Sigma 2.00}}(Mg_{1.16}Fe_{0.39}F_{0.01}Cl_{0.10})_{\Sigma 5.01} \\ T(\underbrace{Si_{7.26}Al_{0.74}}_{E(OH)_{0.39}F_{0.01}Cl_{0.10})_{\Sigma 5.01}}(Mg_{1.16}Fe_{0.39}F_{0.01}Cl_{0.10})_{\Sigma 5.01} \\ T(\underbrace{Si_{7.26}Al_{0.74}}_{E(OH)_{0.39}F_{0.01}Cl_{0.10})_{\Sigma 5.01}}(Mg_{1.16}Fe_{0.39}F_{0.10}Cl_{0.10})_{\Sigma 5.01} \\ T(\underbrace{Si_{7.26}Al_{0.14}}_{E(OH)_{0.14}F_{0.14}Cl_{0.10}}(Mg_{1.16}Fe_{0.14}Fe_{0$ 

significantly with increasing Fe content, a feature which might be used for preliminary mineral petrographic characterization from Moreover, changes in the mean bond lengths measured in the strip of octahedra are in accord with changes in the composition of the samples. The  ${}^{C}Al_{1}{}^{C}Fe_{1}^{3+}$  exchange (occurring at the M(2)site) and the W(OH)<sub>-1</sub>WF<sub>1</sub> exchange (occurring at the O(3) site coordinated by two M(1) and one M(3)cations) describe the transition from magnesiohornblende to magnesio-ferri-fluoro-hornblende, and lead to an increase in the < M(2)-O> distance and a decrease in both < M(1)-O> and < M(3)-O> distances. In contrast, the  ${}^{C}Mg_{-1}{}^{C}Fe_{1}^{2+}$  exchange (occurring at all the M(1,3) sites) and the  ${}^{\rm C}{\rm Al}_{-1}{}^{\rm C}{\rm Fe}_1^{3+}$ exchange (occurring at the M(2) site) describe the transition from magnesio-hornblende to ferro-ferrihornblende and increases all the  $\langle M(1,2,3) - O \rangle$ distances. All these changes are consistent with the differences in ionic radii (Shannon, 1976) of the constituent cations.

The extension of the double chain of tetrahedra along c, as measured by the O(5)–O(6)–O(5) angle, is lower in magnesio-hornblende than in ferro-ferrihornblende; the more extended double chain of tetrahedra promotes linkage with the strip of octahedra where Mg and Al are replaced, respectively, by the larger Fe<sup>2+</sup> and Fe<sup>3+</sup> cations. However, the O(5)–O(6)–O(5) angle in magnesio-ferrifluoro-hornblende (where the <sup>T</sup>Al content is 0.37 apfu larger) is very similar to that in ferro-ferri-hornblende (Oberti *et al.*, 2016*b*).

#### Acknowledgements

The authors wish to thank Fernando Cámara and Peter Leverett for their kind and useful reviews.

## Supplementary material

To view supplementary material for this article, please visit https://doi.org/10.1180/minmag.2017.081.093

# References

- Bartelmehs, K.L., Bloss, F.D., Downs, R.T. and Birch, J.B. (1992) EXCALIBR II. Zeitschrift für Kristallographie, 199, 185–196.
- Bruker (2003) SAINT Software Reference Manual. Version 6. Bruker AXS Inc., Madison, Wisconsin, USA.
- Cannillo, E., Germani, G. and Mazzi, F. (1983) New crystallographic software for Philips PW1100 single

- crystal diffractometer. CNR Centro di Studio per la Cristallografia, Internal Report 2.
- Garzanti, E., Andò, S., Vezzoli, G., Lustrino, M., Boni, M. and Vermeesch, P. (2012) Petrology of the Namib Sand Sea: Long-distance transport and compositional variability in the wind-displaced Orange Delta. *Earth-Science Reviews*, 112, 173–189.
- Hawthorne, F.C., Ungaretti, L. and Oberti, R. (1995) Site populations in minerals: terminology and presentation of results of crystal-structure refinement. *Canadian Mineralogist*, 33, 907–911.
- Hawthorne, F.C., Oberti, R., Harlow, G.E., Maresch, W.V., Martin, R.F., Schumacher, J.C., and Welch, M.D. (2012) Nomenclature of the amphibole supergroup. *American Mineralogist*, 97, 2031–2048.
- Mandarino, J.A. (2007) The Gladstone-Dale compatibility of minerals and its use in selecting mineral species for further studying. *Canadian Mineralogist*, 45, 1307–1324.
- Oberti, R., Ungaretti, L., Cannillo, E. and Hawthorne, F.C. (1992) The behaviour of Ti in amphiboles: I. Four- and six-coordinated Ti in richterites. *European Journal of Mineralogy*, 4, 425–439.
- Oberti, R., Hawthorne, F.C., Ungaretti, L. and Cannillo, E. (1995) <sup>[6]</sup>Al disorder in amphiboles from mantle peridotite. *Canadian Mineralogist*, 33, 867–878.
- Oberti, R., Hawthorne, F.C., Cannillo, E. and Cámara, F. (2007) Long-range order in amphiboles. Pp. 125–172 in: Amphiboles: Crystal Chemistry, Occurrence and Health Issues (F.C. Hawthorne, R. Oberti, G. Della Ventura and A. Mottana, editors). Reviews in Mineralogy & Geochemistry, 67. Mineralogical Society of America and the Geochemical Society, Chantilly, Virginia, USA.
- Oberti, R., Boiocchi, M., Hawthorne, F.C., Ball, N.A. and Chiappino, L. (2016a) Magnesio-ferri-fluorohornblende from Portoscuso, Sardinia, Italy: description of a newly approved member of the amphibole supergroup. *Mineralogical Magazine*, 80, 269–275.
- Oberti, R., Boiocchi, M., Hawthorne, F.C., Ball, N.A., Cámara, F., Pagano, R. and Pagano, A. (2016b) Ferroferri-hornblende from the Traversella Mine (Ivrea, Italy): occurrence, mineral description and crystalchemistry. *Mineralogical Magazine*, 80, 1233–1242.
- Oberti, R., Della Ventura, G., Boiocchi, M., Zanetti, A. and Hawthorne, F.C. (2017) New data on the crystal-chemistry of oxo-mangani-leakeite and mangano-mangani-ungarettiite from the Hoskins mine and their impossible solid-solution An XRD and FTIR study. *Mineralogical Magazine*, 81, 707–722.
- Robinson, K., Gibbs, G.V. and Ribbe, P.H. (1971) Quadratic elongation: a quantitative measure of distortion in coordination polyhedra. *Science*, 172, 567–570.
- Shannon, R.D. (1976) Revised effective ionic radii and systematic studies of interatomic distances in halides and chalcogenides. Acta Crystallographica, A32, 751–767.