

# Folvikite, $\text{Sb}^{5+}\text{Mn}^{3+}(\text{Mg}, \text{Mn}^{2+})_{10}\text{O}_8(\text{BO}_3)_4$ , a new oxyborate mineral from the Kitteln mine, Nordmark ore district, Värmland, Sweden: description and crystal structure

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## ABSTRACT

Folvikite,  $\text{Sb}^{5+}\text{Mn}^{3+}(\text{Mg}, \text{Mn}^{2+})_{10}\text{O}_8(\text{BO}_3)_4$ , is a new oxyborate mineral from the Kitteln mine, Värmland, Sweden, where it occurs as a primary skarn mineral embedded in calcite. It forms striated prismatic crystals up to 0.3 mm, and is black to dark reddish-brown with submetallic lustre and a reddish-brown streak. It is brittle, has a Mohs hardness of 6, and the calculated density is 4.14 g/cm<sup>3</sup>. Folvikite is biaxial with indeterminate optic sign due to pervasive twinning. The optic axial angle is 68.9(4)°. Refractive indices were not measured; the calculated mean refractive index is 1.85. Strong pleochroism was observed in plane-polarized light: AB = brown (intermediate), OB = dark brown (maximum) and ON = honey brown (minimum). Folvikite is monoclinic, space group *P*2<sub>1</sub>, *a* = 5.3767(10), *b* = 6.2108(10), *c* = 10.9389(18) Å, β = 94.399(9)°, *V* = 364.22(16) Å<sup>3</sup> and *Z* = 1. Chemical analysis by electron microprobe gave Sb<sub>2</sub>O<sub>5</sub> 18.15, MgO 24.11, MnO 29.73, Mn<sub>2</sub>O<sub>3</sub> 11.62, Al<sub>2</sub>O<sub>3</sub> 0.27, Fe<sub>2</sub>O<sub>3</sub> 0.45, B<sub>2</sub>O<sub>3</sub> 15.27, sum 99.60 wt.%. The B<sub>2</sub>O<sub>3</sub> content was assigned as B = 4 apfu and the Mn<sub>2</sub>O<sub>3</sub> / (MnO + Mn<sub>2</sub>O<sub>3</sub>) ratio was determined from the crystal structure. The empirical formula was normalized on the basis of 20 anions pfu: (Sb<sup>5+</sup><sub>1.02</sub>Mn<sup>3+</sup><sub>1.34</sub>Al<sub>0.05</sub>Fe<sup>3+</sup><sub>0.05</sub>Mg<sub>5.46</sub>Mn<sup>2+</sup><sub>3.82</sub>□<sub>0.26</sub>)<sub>Σ12</sub>O<sub>8</sub>(BO<sub>3</sub>)<sub>4</sub>. A simplified formula may be written as  $\text{Sb}^{5+}\text{Mn}^{3+}(\text{Mg}, \text{Mn}^{2+})_{10}\text{O}_8(\text{BO}_3)_4$  with *Z* = 1. The crystal structure was solved by direct methods and refined to an *R*<sub>1</sub> index of 4.1%. Folvikite is a member of the (3 Å) zigzag wallpaper-borate structures in which chains of edge-sharing octahedra extend along the *c* axis and are cross-linked by BO<sub>3</sub> groups. There are five *X* sites partly occupied by Mn<sup>2+</sup> > Mg, one octahedrally coordinated *M*-site occupied by Sb<sup>5+</sup> > Mg, two *M* sites occupied by Mg ≥ Mn > Sb<sup>5+</sup>, two *M* sites occupied by Mn<sup>3+</sup> > Mn<sup>2+</sup>, two *M* sites occupied by Mg > Mn<sup>2+</sup>, and one *M*-site occupied by Mg > □; plus two [3]-coordinated *B* sites occupied by B. As with the other zigzag borates, the polyhedra are arranged in F-walls, C-walls and S-columns.

**KEYWORDS:** folvikite, new mineral, chemical analysis, crystal structure, infrared spectrum, Kitteln mine, Värmland, Sweden.

## Introduction

FOLVIKITE is a new Sb-bearing oxyborate mineral of the ‘pinakiolite group’ of Strunz and Nickel (2001). It was found at the Kitteln mine of the Nordmark ore field, north of Filipstad, Värmland, Sweden

(59°49′59″N, 14°05′59″E). The Kitteln mine is also the type locality for the Sb-bearing oxyborate blatterite (Raade *et al.*, 1988; Cooper and Hawthorne, 1998).

The mineral is named after Harald O. Folvik (born 1941). Mr. Folvik is a prominent Norwegian amateur mineralogist who has a particular interest in Långban-type deposits and is a member of The Långban Society. The mineral described here was noticed initially by him, and he collected the specimens that

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were used in this work. Both the mineral and the name have been approved by the Commission on New Minerals, Nomenclature and Classification of the International Mineralogical Association (IMA2016-026). The polished mount used for electron probe microanalysis is kept at the Natural History Museum, University of Oslo, Norway (catalogue number 43574). The crystal used for the structural study, powdered material used for IR spectroscopy and subsequent powder X-ray diffraction and the rock piece that yielded this material were deposited at the Canadian Museum of Nature, Ottawa (catalogue number CMNMC 87087).

## Occurrence

The mines of the Nordmark ore field (Magnusson, 1929) were essentially worked for iron ore. Manganese ores were produced from two small mines: Östra Mossgruvan and Kitteln. The manganese ore at Östra Mossgruvan, originally mined for iron, was discovered in 1876 and consisted mainly of hausmannite with grains of manganosite (Sjögren, 1876, 1878). Production of manganese ore at the Kitteln mine, also at first mined for iron, began in 1887, and the ore was similar to that from Östra Mossgruvan (Sjögren, 1887).

The folvikite-bearing rock material in the Kitteln mine, as well as the blatterite-bearing material, was evidently dumped there from elsewhere. In the description of blatterite, Brattforsgruvan was tentatively suggested as the most probable source, based on the fact that katoptrite occurs with blatterite at Kitteln, and Brattforsgruvan is the type locality of that mineral. However, manganese ore has not, as far as we know, been reported from Brattforsgruvan, and Östra Mossgruvan is more likely to be the source. The distance between Kitteln and Östra Mossgruvan is only ~130–140 m. It is also conceivable that the backfill originates from the Kitteln mine itself. Considering this uncertainty, the Kitteln mine, where the material was collected, should be regarded as the type locality of both blatterite and folvikite.

## Associated minerals

Folvikite occurs embedded in calcite. Other minerals identified in the assemblage are hausmannite, tegengrenite and a Mg-dominant analogue of sonolite or jerrygibbsite,  $(\text{Mg}, \text{Mn})_9(\text{SiO}_4)_4(\text{OH})_2$ ; the latter species may occur as inclusions in folvikite. Blatterite and folvikite do not occur together, blatterite being confined to manganosite-bearing ore.

## Physical properties

Folvikite forms strongly lustrous, elongate prismatic crystals up to 0.3 mm long. They are commonly striated along their length due to twinning by  $180^\circ$  rotation about  $c^*$ ; an aggregate, 0.1 mm wide, of five parallel individuals, was reported by Raade and Folvik (2013). Crystals are elongated along [010] with indeterminate interpenetrating  $\{h0l\}$  forms (Fig. 1). Folvikite is black to dark reddish-brown with a reddish-brown streak. It has a submetallic lustre and the diaphaneity is generally opaque except at thin edges where the colour is dark red-brown. Cleavage and parting were not observed, folvikite is brittle and has conchoidal fracture. Mohs hardness is 6, and folvikite does not luminesce in ultraviolet light. The calculated density is  $4.14 \text{ g/cm}^3$  (from the empirical formula). Folvikite is biaxial with indeterminate optic sign due to pervasive twinning. The optical axial angle was measured with a spindle stage using the program *Excalibr2* (Bartelmehs *et al.*, 1992) and is  $68.9(4)^\circ$ . Refractive indices were not measured; the calculated mean refractive index is 1.85. The long axis of the crystal is parallel to the optic normal and hence  $Y=b$ . Strong pleochroism was observed in plane-polarized light: AB = brown (intermediate), OB = dark brown (maximum) and ON = honey brown (minimum).

## Chemical composition

Crystals were analysed with a Cameca SX100 electron microprobe (EMPA) operated in wavelength-dispersive mode at 15 kV and 15 nA and using a beam diameter of 1  $\mu\text{m}$ . Concentrations of Ca and Ti were below the detection limits of EMPA.



FIG. 1. Back-scattered electron image of crystals of folvikite; note the interpenetrating forms.

The following standards were used:  $\text{Sb}_2\text{S}_3$  (Sb),  $\text{MgO}$  (Mg), pyrophanite (Mn),  $\text{Al}_2\text{O}_3$  (Al) and Fe metal (Fe). Boron was analysed by electron counting during crystal-structure refinement (Hawthorne and Grice, 1990), and its presence in the structure was also confirmed by infrared spectroscopy (see below). Lithium is below detection using LA-ICP-MS (laser ablation inductively coupled plasma mass spectrometry). Table 1 gives the chemical composition (mean of seventeen points from three grains). The empirical formula unit, calculated to give a cation sum (excluding B) of 11.74 apfu with  $B = 4$  apfu (atoms per formula unit) is:  $(\text{Sb}_{1.02}^{5+}\text{Mn}_{1.34}^{3+}\text{Al}_{0.05}\text{Fe}_{0.05}^{3+}\text{Mg}_{5.46}\text{Mn}_{3.82}^{2+}\square_{0.26})_{\Sigma 12}\text{O}_8(\text{BO}_3)_4$ ; the simplified formula is  $\text{Sb}^{5+}\text{Mn}^{3+}(\text{Mg},\text{Mn}^{2+})_{10}\text{O}_8(\text{BO}_3)_4$  ( $Z = 1$ ) (see later discussion for justification of this procedure).

### Infrared spectroscopy

The Fourier-transform infrared (FTIR) spectrum was collected using a Bruker Hyperion 2000 IR microscope equipped with a liquid-nitrogen-cooled mercury-cadmium-telluride detector. The spectrum in the range  $4000\text{--}650\text{ cm}^{-1}$  was obtained by averaging 100 scans with a resolution of  $4\text{ cm}^{-1}$ . The infrared spectrum of folvikite is shown in Fig. 2. The spectrum shows no absorption bands in the OH-stretching region ( $4000\text{--}3000\text{ cm}^{-1}$ ), which is in accord with the absence of OH groups and  $\text{H}_2\text{O}$  in the structure. Absorption bands in the region  $1500\text{--}900\text{ cm}^{-1}$  (with strong bands at 1335, 1275 and  $1240\text{ cm}^{-1}$ , and medium to low intensity bands at 1124, 948, 934 and  $906\text{ cm}^{-1}$ ) may be

assigned to B–O stretching vibrations of the borate group. Bands at  $723$  and  $666\text{ cm}^{-1}$  are due to bending modes of the borate group.

### Powder X-ray diffraction

The powder-diffraction pattern for folvikite was recorded at the Canadian Museum of Nature using a Bruker D8 Discover micro-powder diffractometer with a Hi-Star multi-wire 2D detector at  $12\text{ cm}$ , calibrated following Rowe (2009) and using a pseudo-Gandolfi approach on a powdered sample of folvikite that was crushed using a diamond micro-anvil. The unit-cell dimensions were refined using the program *UNITCELL* of Holland and Redfern (1995). The powder pattern is given in Table 2, together with the refined cell-dimensions which are in close accord with the values derived by single-crystal X-ray diffraction.

### Crystal structure: data collection, structure solution and refinement

A single (twinned) crystal ( $5\text{ }\mu\text{m} \times 6\text{ }\mu\text{m} \times 40\text{ }\mu\text{m}$ ) was attached to a tapered glass fibre and mounted on a Bruker D8 three-circle diffractometer equipped with a rotating anode generator ( $\text{MoK}\alpha$  X-radiation), multilayer optics and an APEX-II detector. In excess of a Ewald sphere of data was collected to  $60^\circ 2\theta$  using  $26\text{ s}$  per  $0.2^\circ$  frame with a crystal-detector distance of  $8\text{ cm}$ . A preliminary evaluation of the diffraction pattern at a crystal-detector distance of  $5\text{ cm}$  showed that the crystal contained a significant non-merohedral twin component [twin law:  $180^\circ$  rotation about  $c^*$ ], and the detector was moved back to  $8\text{ cm}$  to further enhance resolution of the partly overlapping twin component. Twin integration gave 30,684 total reflections, with 11,869 [component 1], 11,935 [component 2] and 6880 [both components] constituent reflections. The reflections were averaged and merged [ $R_{\text{int}} = 5.8\%$ ] to give 1166 reflections (single reflections from the primary domain, plus composites involving both domains) for structure (twin) refinement. The unit-cell dimensions were obtained by least-squares refinement of 7131 reflections with  $I > 10\sigma I$ .

An oriented ‘precession’ slice ( $h0l$ ) was constructed from the raw data frames and is shown in Fig. 3; the diffraction spots from the primary domain (green  $h0l$  labelling) are related to the diffraction spots from the twin domain (red  $h0l$  labelling) by a  $180^\circ$  rotation about  $c^*$ . Note the different inter-spot spacing along  $c^*$  for different

TABLE 1. Chemical composition (wt.%) of folvikite.

Constituent	Wt.%	Range
$\text{Sb}_2\text{O}_5$	18.15(32)	17.31–18.55
$\text{MgO}$	24.11(66)	22.79–25.06
$\text{MnO}^{\text{tot}}$	40.17(89)	39.04–41.64
$\text{MnO}$	29.73	
$\text{Mn}_2\text{O}_3^*$	11.62	
$\text{Al}_2\text{O}_3$	0.27(2)	0.24–0.31
$\text{Fe}_2\text{O}_3$	0.45(6)	0.36–0.59
$\text{B}_2\text{O}_3^{**}$	15.27	
Total	99.60	

Below detection: Ca and Ti (EMPA); Li (LA-ICP-MS);

\*calculated to give a cation sum (excluding  $\text{B}^{3+}$ ) of 11.74 apfu;

\*\*calculated to give  $\text{B}^{3+} = 4$  apfu.

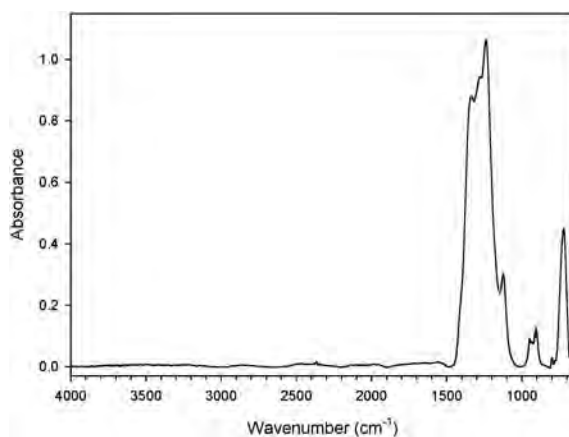


FIG. 2. The FTIR spectrum of folvikite.

TABLE 2. X-ray powder diffraction pattern\*\* of folvikite.

$I_{\text{obs}}$ %	$I_{\text{SXL}}$ %	$d_{\text{SXL}}$ Å	$d_{\text{obs}}$ Å	$d_{\text{calc}}$ Å	$h k l$
<b>100</b>	100	5.453	<b>5.450</b>	5.450	0 0 2
	34	5.361			1 0 0
<b>46</b>	46	4.964	<b>4.973</b>	4.972	$\bar{1}$ 0 1
12	14	3.684	3.688	3.687	1 0 2
32	34	3.106	3.112	3.110	0 2 0
10	11	2.907	2.907 *		1 0 3
<b>82</b>	95	2.727	<b>2.725</b>	2.725	0 0 4
	24	2.699			0 2 2
<b>91</b>	56	2.586	<b>2.591</b>	2.590	1 2 1
31	43	2.482	2.486	2.486	$\bar{2}$ 0 2
15	12	2.448	2.450	2.451	$\bar{1}$ 2 2
24	15	2.375	2.375	2.377	1 2 2
9	10	2.336	2.338	2.339	2 0 2
8	7	2.202	2.201	2.204	$\bar{1}$ 2 3
13	22	2.181	2.183	2.180	0 0 5
12	7	2.122	2.124	2.124	1 2 3
<b>43</b>	35	2.029	<b>2.033</b>	2.033	2 2 0
	13	2.016			$\bar{2}$ 2 1
5	16	1.989	1.991	1.991	$\bar{2}$ 0 4
3	3	1.939	1.940	1.942	$\bar{2}$ 2 2
15	10	1.878	1.880	1.879	1 2 4
11	9	1.675	1.672 *		$\bar{2}$ 2 4
15	15	1.663	1.662	1.663	1 2 5
15	15	1.584	1.585	1.586	$\bar{2}$ 2 4
	18	1.561			2 0 6
<b>37</b>	42	1.553	<b>1.552 *</b>		0 4 0
14	17	1.533	1.532	1.533	$\bar{1}$ 2 6

$a = 5.388(2)$ ,  $b = 6.220(2)$ ,  $c = 10.933(3)$  Å,  $\beta = 94.40(3)^\circ$  and  $V = 365.27(15)$  Å<sup>3</sup>

\*Excluded in least-squares refinement; \*\*the subscript SXL denotes single-crystal diffraction data.

The six strongest lines are given in bold.

values of  $H$ ; complete overlap occurs for  $(00l)$  reflections, and nearly so for  $(h0l)$  reflections with  $H = 3$  and  $-3$ . All diffraction maxima from the X-ray crystal can be indexed on the small monoclinic cell ( $5.38 \times 6.21 \times 10.94 \text{ \AA}$ ) with the inclusion of the twin law. Structure (twin) refinement from 1166 reflections (including 607 composites) gave a final  $R_1$  value of 4.1% (for 1070 observed reflections,  $|F_o| > 4\sigma F$ ). The Flack parameter refined to  $-0.03$  (16), indicating a correct absolute structural configuration. The twin volume fraction (i.e. twin contribution to composites) refined to 0.417(2), suggesting a very significant twin fraction (as is indicated on the precession slice in Fig. 3). Miscellaneous information concerning structure solution and refinement is listed in Table 3. Atom positions and displacement parameters are given in Table 4, selected interatomic distances in Table 5, site-scattering values and site assignments in Table 6 and a bond-valence calculation in Table 7. The crystallographic information files have been deposited with the Principal Editor of *Mineralogical Magazine* and are available as Supplementary material (see below).

### Crystal structure: site assignment and formula normalization

The crystal structure of folvikite contains both positionally ordered  $M$ -cations (octahedral coordination), positionally disordered  $X$ -cations (coordination numbers [4], [5] and [6]), and ordered  $\text{BO}_3$  groups. There are eight  $M$ -cation sites and five  $X$ -cation sites, and all give refined site-scattering values (Hawthorne *et al.*, 1995) that suggest multiple occupants at each of these sites. There are primarily five constituents to assign over the thirteen ( $M + X$ ) sites:  $\text{Sb}^{5+}$ ,  $\text{Mg}^{2+}$ ,  $\text{Mn}^{2+}$ ,  $\text{Mn}^{3+}$  and  $\square$  (vacancy); we combined the minor  $\text{Al}^{3+}$  and  $\text{Fe}^{3+}$  contents from the chemical data with  $\text{Mn}^{3+}$  on the basis of their similar charge and radii (Shannon, 1976).

As the structure of folvikite contains the same architectural elements as in other zigzag borates (i.e. F- and C-walls; and S-columns; Cooper and Hawthorne, 1998), we use many of the same general crystal-chemical arguments to help assign cations in the folvikite structure. The  $M$  sites in folvikite have refined site-scattering values spanning the range 8.9–41.6 epfu (electrons per formula unit), and a vacancy ( $\square$ ) must be associated with at least one of these sites, as Mg (12 electrons) is the weakest X-ray scattering species present. With a

known vacancy present for at least one  $M$ -site, and additionally at the positionally disordered  $X$  sites, the overall cation sum for folvikite is not fixed. An additional variable involves the partitioning of Mn as  $\text{Mn}^{2+}$  and  $\text{Mn}^{3+}$ . The most feasible approach to normalizing the chemical data therefore involves an anion-normalization scheme [i.e.  $(\text{BO}_3)_4\text{O}_8 = 28^-$ ], with an assessment of both the vacancy content and the  $\text{Mn}^{2+}/\text{Mn}^{3+}$  ratio via crystal-chemical arguments from the structure refinement. The  $M(1)$  site has the largest refined site-scattering value (41.6 epfu) and the shortest mean bond length ( $\langle M(1)-\text{O} \rangle = 2.019 \text{ \AA}$ ) of all the  $M$  sites and must contain significant  $\text{Sb}^{5+}$  (Table 6). Note that there is no  $\text{Sb}^{3+}$  present in folvikite, as  $\text{Sb}^{3+}$  is a lone-pair stereoactive cation that has very different stereochemistry that we do not observe here (Mills *et al.*, 2009). The  $M(2)$  site has the lowest refined site-scattering value (8.9 epfu) and a  $\langle M(2)-\text{O} \rangle$  distance of 2.119  $\text{\AA}$ ; without knowing the constituent size of a vacancy, or whether there is more than one type of cation present, the simplest assignment to be made for the  $M(2)$  site involves Mg + vacancy (Table 6). The next two smallest  $M$ -octahedra involve the  $M(4)$  and  $M(5)$  sites with  $\langle M(4)-\text{O} \rangle = 2.065 \text{ \AA}$  and  $\langle M(5)-\text{O} \rangle = 2.071 \text{ \AA}$ ; the refined site-scattering values for these two sites ( $\sim 23$  epfu), in combination with the observed  $\langle M-\text{O} \rangle$  distances, suggest dominant  $\text{Mn}^{3+}$ . In addition, by comparison with other zigzag borate structures, we expect these edge-sharing [ $= M(4) = M(5) =$ ] strips of octahedra within the C-wall to be occupied by Jahn-Teller-distorted  $\text{Mn}^{3+}$  octahedra. Most of the pinakiolite-group minerals contain [4 + 2]-distorted  $\text{Mn}^{3+}$ -bearing octahedra with four short and two long bonds; however in folvikite, the  $\text{Mn}^{3+}$  octahedra show [2 + 2 + 2] distortion (i.e. two short, two intermediate and two long bonds) and are occupied by  $\sim 33\%$  other cations. This situation is analogous to that of the  $M(4)$  octahedron in fredrikssonite (Burns *et al.*, 1994) in which a similar [2 + 2 + 2] distortion occurs and the  $M(4)$  site is occupied primarily by  $\text{Mn}^{3+}$  with  $\sim 22\%$  other cations. The cations Mg,  $\text{Mn}^{2+}$  and minor  $\text{Sb}^{5+}$  were assigned to the remaining  $M$  sites [ $M(3)$ ,  $M(6)$ ,  $M(7)$  and  $M(8)$ ] in folvikite in accord with the relative site-scattering values and observed bond lengths. The positionally disordered  $X$  sites adjacent to the Jahn-Teller-distorted  $\text{Mn}^{3+}$  octahedral strips of the C-wall are considered to be occupied predominately by  $\text{Mn}^{2+}$ , with subordinate Mg. There are five  $X$  sites in total, and we propose maximum occupancy for this region of the structure (i.e. 2 apfu), whereby close  $X-X'$  contacts preclude local mutual occupancy of many  $X-X'$  pairs.

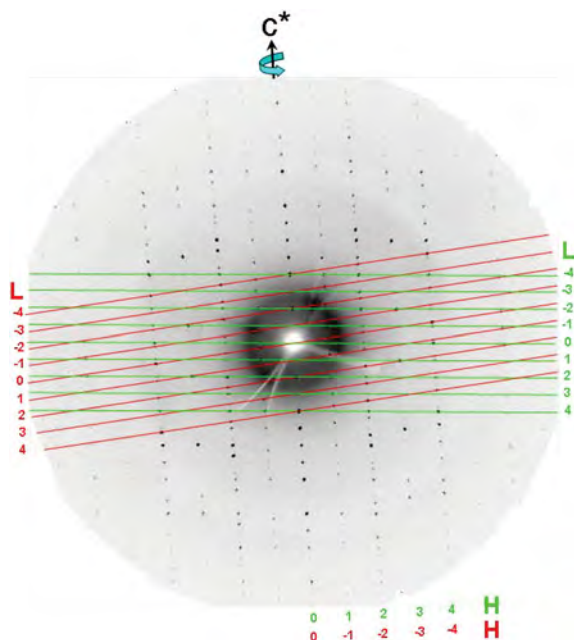


FIG. 3. The  $(h0l)$  precession layer for folvikite generated from the single-crystal X-ray diffraction data. Twin law =  $180^\circ$  rotation about  $c^*$ . Primary-domain indexing in green and twin-component indexing in red.

The total refined site-scattering at all  $(M+X)$  sites is 247.3 epfu [ $Z=1$ ], in good agreement with the scattering value of 248.7 epfu calculated for the

TABLE 3. Miscellaneous information concerning data collection and structure refinement for folvikite.

$a$ (Å)	5.3767(10)
$b$	6.2108(10)
$c$	10.9389(18)
$\beta$ ( $^\circ$ )	94.399(9) $^\circ$
$V$ (Å <sup>3</sup> )	364.22(16)
Space group	$P2_1$
$Z$	4
$D_{\text{calc}}$ (g/cm <sup>3</sup> )	4.14
Crystal size ( $\mu\text{m}$ )	$5 \times 6 \times 40$
Radiation	MoK $\alpha$
No. of reflections	30,684
No. unique reflections	1166
$R_{\text{int}}$ %	5.8
No. with $(F_o > 4\sigma F)$	1070
$R_1$ %	4.1
$wR_2$ %	8.0
Goof	1.103
Parameters/restraints	127/1

$R_1 = \Sigma(|F_o| - |F_c|) / \Sigma|F_c|$   
 $wR_2 = [\Sigma w(F_o^2 - F_c^2)^2 / \Sigma w(F_o^2)^2]^{1/2}$ ,  $w = 1 / [\sigma^2(F_o^2) + (0.0122 P)^2 + 3.90 P]$  where  $P = (\max(F_o^2, 0) + 2F_c^2) / 3$

chemical formula derived from electron-microprobe analysis. The chemical data were normalized to 20  $\text{O}^{2-}$  anions, with 4 B apfu, i.e.  $[\text{O}_8(\text{BO}_3)_4] = 28^-$ , with a  $\text{Mn}^{3+}/(\text{Mn}^{3+} + \text{Mn}^{2+})$  value adjusted to give an  $[M+X]$  cation sum of 11.74 apfu [i.e.  $12.00 - 0.26^{M(2)}\square$ ]. The constituent radii for the  $M$ -site assignments are plotted against the observed  $\langle M\text{-O} \rangle$  distances in Fig. 4, and a linear regression gives  $\langle M\text{-O} \rangle = 0.834 (r) + 1.497 \text{ \AA}$ . A similar regression line constructed for the octahedrally coordinated sites in blatterite gave  $\langle M\text{-O} \rangle = 0.8325 (r) + 1.5060$  (Cooper and Hawthorne, 1998), and the congruency between the two relations supports the site assignments for these two homologous structures. Finally, the incident bond-valence sums at the  $M$  sites (Table 7) are in close accord with the aggregate formal charge of the constituent cations at these sites, with a mean deviation of 0.07. Moreover, the incident bond-valence sums at the  $B$  sites (Table 7) are in accord with their occupancy by  $\text{B}^{3+}$ .

Crystal structure: description

As the structure of folvikite is related closely to the structure of pinakiolite, we describe its relation to its closest structural (and chemical)



TABLE 4. Atom positions and equivalent isotropic-displacement parameters ( $\text{\AA}^2$ ) for folvikite.

Site	<i>x</i>	<i>y</i>	<i>z</i>	$U_{\text{eq}}$	$U^{11}$	$U^{22}$	$U^{33}$	$U^{23}$	$U^{13}$	$U^{12}$
<i>M</i> (1)	½	0.9064(4)	0	0.0038(3)	0.0018(4)	0.0048(4)	0.0047(5)	0	0.0006(4)	0
<i>M</i> (2)	0	0.168(2)	0	0.014(4)	0.017(6)	0.002(4)	0.024(7)	0	−0.001(6)	0
<i>M</i> (3)	½	0.4093(13)	0	0.0128(7)	0.0118(10)	0.0111(10)	0.0155(12)	0	0.0011(10)	0
<i>M</i> (4)	0	0.4069(11)	½	0.0093(6)	0.0093(9)	0.0077(9)	0.0103(10)	0	−0.0025(7)	0
<i>M</i> (5)	½	0.4048(10)	½	0.0100(6)	0.0107(9)	0.0071(9)	0.0127(10)	0	0.0036(7)	0
<i>M</i> (6)	0	0.6521(12)	0	0.0101(15)	0.007(2)	0.017(2)	0.007(2)	0	0.003(2)	0
<i>M</i> (7)	0.7176(8)	0.6574(7)	0.2767(2)	0.0078(9)	0.0071(19)	0.0082(14)	0.0083(16)	0.001(3)	0.0018(17)	−0.003(3)
<i>M</i> (8)	0.7166(8)	0.1597(7)	0.2760(2)	0.0099(9)	0.0112(19)	0.0106(14)	0.0078(15)	0.000(3)	0.0011(17)	−0.003(3)
<i>X</i> (1)	½	0.911(3)	½	0.007(2)						
<i>X</i> (2)	0.401(2)	0.907(4)	0.4834(7)	0.008(3)						
<i>X</i> (3)	0.241(2)	0.904(4)	0.4524(7)	0.017(3)						
<i>X</i> (4)	0	0.912(3)	½	0.007(2)						
<i>X</i> (5)	0.098(3)	0.908(5)	0.4866(9)	0.008(4)						
<i>B</i> (1)	0.207(2)	0.409(5)	0.2295(8)	0.0080(14)						
<i>B</i> (2)	0.212(2)	0.910(6)	0.2387(9)	0.0130(17)						
O(1)	0.9946(11)	0.406(3)	0.2926(6)	0.0111(13)						
O(2)	0.4415(11)	0.412(3)	0.2905(5)	0.0103(13)						
O(3)	0.1826(13)	0.407(3)	0.1018(5)	0.0076(11)						
O(4)	0.9950(11)	0.908(3)	0.2966(6)	0.0111(12)						
O(5)	0.4487(11)	0.912(3)	0.2978(6)	0.0123(13)						
O(6)	0.1946(14)	0.904(3)	0.1069(5)	0.0097(12)						
O(7)	0.677(2)	0.1375(18)	0.0909(8)	0.009(2)						
O(8)	0.685(2)	0.663(2)	0.0885(8)	0.009(2)						
O(9)	0.7463(19)	0.6081(14)	0.4709(7)	0.0060(16)						
O(10)	0.7446(19)	0.1595(19)	0.4727(7)	0.0105(18)						

TABLE 5. Selected interatomic distances (Å) in folvikite.

$M(1)-O(6) \times 2$	2.088(7)	$X(1)-O(5) \times 2$	2.208(6)
$M(1)-O(7) \times 2$	1.952(11)	$X(1)-O(9) \times 2$	2.336(18)
$M(1)-O(8) \times 2$	2.017(12)	$X(1)-O(10) \times 2$	2.064(19)
$\langle M(1)-O \rangle$	2.019	$\langle X(1)-O \rangle$	2.203
$M(2)-O(3) \times 2$	2.056(17)	$X(2)-O(5)$	2.067(10)
$M(2)-O(6) \times 2$	2.228(18)	$X(2)-O(5)$	2.467(11)
$M(2)-O(7) \times 2$	2.074(12)	$X(2)-O(9)$	2.09(2)
$\langle M(2)-O \rangle$	2.119	$X(2)-O(10)$	1.83(2)
$M(3)-O(3) \times 2$	2.108(7)	$X(2)-O(10)$	2.43(2)
$M(3)-O(7) \times 2$	2.145(14)	$\langle X(2)-O \rangle$	2.18
$M(3)-O(8) \times 2$	2.062(14)	$X(3)-O(4)$	2.075(11)
$\langle M(3)-O \rangle$	2.105	$X(3)-O(5)$	2.098(11)
$M(4)-O(1) \times 2$	2.267(6)	$X(3)-O(9)$	2.02(2)
$M(4)-O(9) \times 2$	1.860(10)	$X(3)-O(10)$	1.78(3)
$M(4)-O(10) \times 2$	2.067(11)	$\langle X(3)-O \rangle$	1.99
$\langle M(4)-O \rangle$	2.065	$X(4)-O(4) \times 2$	2.223(6)
$M(5)-O(2) \times 2$	2.290(6)	$X(4)-O(9) \times 2$	2.336(17)
$M(5)-O(9) \times 2$	1.875(10)	$X(4)-O(10) \times 2$	2.068(18)
$M(5)-O(10) \times 2$	2.049(11)	$\langle X(4)-O \rangle$	2.209
$\langle M(5)-O \rangle$	2.071	$X(5)-O(4)$	2.110(11)
$M(6)-O(3) \times 2$	2.088(13)	$X(5)-O(4)$	2.461(12)
$M(6)-O(6) \times 2$	2.174(14)	$X(5)-O(9)$	2.08(3)
$M(6)-O(8) \times 2$	2.016(12)	$X(5)-O(10)$	1.81(3)
$\langle M(6)-O \rangle$	2.093	$X(5)-O(10)$	2.46(2)
$M(7)-O(1)$	2.156(16)	$\langle X(5)-O \rangle$	2.18
$M(7)-O(2)$	2.142(15)	$B(1)-O(1)$	1.377(11)
$M(7)-O(4)$	2.153(16)	$B(1)-O(2)$	1.383(11)
$M(7)-O(5)$	2.167(17)	$B(1)-O(3)$	1.393(10)
$M(7)-O(8)$	2.053(9)	$\langle B(1)-O \rangle$	1.384
$M(7)-O(9)$	2.140(7)	$B(2)-O(4)$	1.371(13)
$\langle M(7)-O \rangle$	2.135	$B(2)-O(5)$	1.382(13)
$M(8)-O(1)$	2.136(16)	$B(2)-O(6)$	1.438(11)
$M(8)-O(2)$	2.168(16)	$\langle B(2)-O \rangle$	1.397
$M(8)-O(4)$	2.166(16)		
$M(8)-O(5)$	2.134(16)		
$M(8)-O(7)$	2.025(9)		
$M(8)-O(10)$	2.145(7)		
$\langle M(8)-O \rangle$	2.129		

relative: Sb-rich pinakiolite from Långban, Sweden (Hansen *et al.*, 1988; Norrestam and Hansen, 1990). Folvikite and Sb-rich pinakiolite share the same style of structural connectivity, but differ in terms of unit cell and space group: folvikite crystallizes in  $P2$  and has a cell volume half that of Sb-rich pinakiolite which crystallizes in  $C2/m$ .

#### F-walls

In the crystal structure of folvikite,  $M(1)$ ,  $M(2)$ ,  $M(3)$  and  $M(6)$  octahedra share edges to form the F-wall (Fig. 5a), whereas in Sb-rich pinakiolite the F-wall is comprised of  $M(3)$  and  $M(6)$  edge-sharing octahedra (Fig. 5b). The largest difference between the F-walls of the two structures involves ordering of  $Sb^{5+}$ : in



TABLE 6. Site-scattering values (epfu) and assigned site-populations (apfu).

Site	Mult. <sup>a</sup>	SFAC <sup>b</sup>	SREF <sup>c</sup>	Sb <sup>5+</sup>	Mn <sup>3+</sup>	Al	Fe <sup>3+</sup>	Mg	Mn <sup>2+</sup>	□	epfu <sup>d</sup>
<i>M</i> (1)	1	Sb,Mg	41.6(3)	0.75				0.25			41.3
<i>M</i> (2)	1	Mg	8.9(4)					0.74		0.26	8.9
<i>M</i> (3)	1	Sb,Mg	23.6(3)	0.16				0.40	0.44		24.0
<i>M</i> (4)	1	Mn	22.8(3)		0.67	0.05		0.08	0.20		23.4
<i>M</i> (5)	1	Mn	23.2(3)		0.67		0.05	0.08	0.20		24.0
<i>M</i> (6)	1	Sb,Mg	18.1(4)	0.11				0.76	0.13		18.0
<i>M</i> (7)	2	Mn,Mg	32.4(5)					1.36	0.64		32.3
<i>M</i> (8)	2	Mn,Mg	33.9(5)					1.24	0.76		33.9
<i>X</i> (1)	1	Mn	9.6(5)*								
<i>X</i> (2)	2	Mn	8.1(6)*								
<i>X</i> (3)	2	Mn	8.4(4)*								42.9
<i>X</i> (4)	1	Mn	10.4(5)*					0.55*	1.45*		
<i>X</i> (5)	2	Mn	6.3(6)*								
Total			247.3	1.02	1.34	0.05	0.05	5.46	3.82	0.26	248.7

<sup>a</sup>Site multiplicity<sup>b</sup>Assigned scattering factors<sup>c</sup>Refined site-scattering values (epfu)<sup>d</sup>Assigned electrons per formula unit [*Z* = 1]\*Total refined site-scattering over all *X* sites = 42.8 epfu attributed to (Mn<sub>1.45</sub><sup>2+</sup>Mg<sub>0.55</sub>)

TABLE 7. Bond-valence (in valence units) table for folvikite.

	O(1)	O(2)	O(3)	O(4)	O(5)	O(6)	O(7)	O(8)	O(9)	O(10)	$\Sigma$
<i>M</i> (1)						0.59 <sup>x2</sup>	0.85 <sup>x2</sup>	0.72 <sup>x2</sup>			4.32
<i>M</i> (2)			0.37 <sup>x2</sup>			0.24 <sup>x2</sup>	0.36 <sup>x2</sup>				1.94
<i>M</i> (3)			0.42 <sup>x2</sup>				0.38 <sup>x2</sup>	0.47 <sup>x2</sup>			2.54
<i>M</i> (4)	0.25 <sup>x2</sup>								0.75 <sup>x2</sup>	0.43 <sup>x2</sup>	2.86
<i>M</i> (5)		0.24 <sup>x2</sup>							0.73 <sup>x2</sup>	0.46 <sup>x2</sup>	2.86
<i>M</i> (6)			0.39 <sup>x2</sup>			0.31 <sup>x2</sup>		0.48 <sup>x2</sup>			2.36
<i>M</i> (7)	0.32	0.33		0.32	0.31			0.42	0.33		2.03
<i>M</i> (8)	0.33	0.31		0.31	0.33		0.46			0.32	2.06
<i>B</i> (1)	0.98	0.97	0.94								2.89
<i>B</i> (2)				1.00	0.97	0.83					2.80
<i>X</i> (1)					<b>0.30</b> <sup>x2</sup>				<b>0.22</b> <sup>x2</sup>	<b>0.45</b> <sup>x2</sup>	1.94
<i>X</i> (2)					0.15				0.41	0.17	2.01
					0.44					0.84	
<i>X</i> (3)				0.43	0.40				0.50	0.96	2.29
<i>X</i> (4)				<b>0.29</b> <sup>x2</sup>					<b>0.22</b> <sup>x2</sup>	<b>0.44</b> <sup>x2</sup>	1.90
<i>X</i> (5)				0.15					0.43	0.15	2.01
				0.39						0.89	
$\Sigma$	1.88	1.85	2.12	<b>1.92</b>	<b>1.91</b>	1.97	2.05	2.09	<b>2.25</b>	<b>2.10</b>	

\*Using the bond-valence parameters of Brown and Altermatt (1985); all bond-valence multipliers (i.e. x2) apply to the summations at the cations; bond valences in bold font are for a simple *X*(1), *X*(4) ordered arrangement.

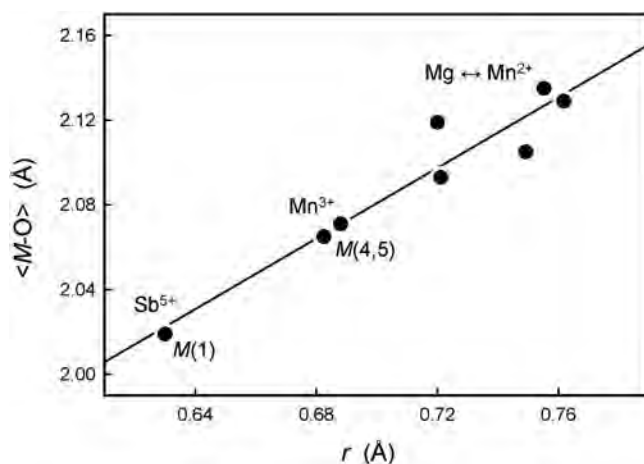


FIG. 4. The constituent radii for the  $M$ -site assignments versus the observed  $\langle M-O \rangle$  distances for folvikite.

Sb-rich pinakiolite, the  $M(3)$  site is occupied by  $Sb_{0.38}^{5+}Mn_{0.32}^{2+}Mg_{0.30}$ , whereas in folvikite, the comparable octahedron strip has  $Sb^{5+}$  segregated differently between  $M(1) = (Sb_{0.75}^{5+}Mg_{0.25})$  and  $M(3) = (Mn_{0.44}^{2+}Mg_{0.40}Sb_{0.16}^{5+})$  (Table 6). This cation ordering at  $M(1)$  and  $M(3)$  in folvikite leads to a mean difference in octahedron size that is large enough to

be visually discernible (Fig. 5a). Many zigzag-borate minerals bearing significant Sb (i.e. folvikite, Sb-rich pinakiolite and blatterite) have so far shown that all Sb is present as  $Sb^{5+}$  that occupies octahedrally coordinated sites within the F-wall (Norrestam and Hansen, 1990; Cooper and Hawthorne, 1998). We note here that the  $Sb_2O_5$  content is essentially the

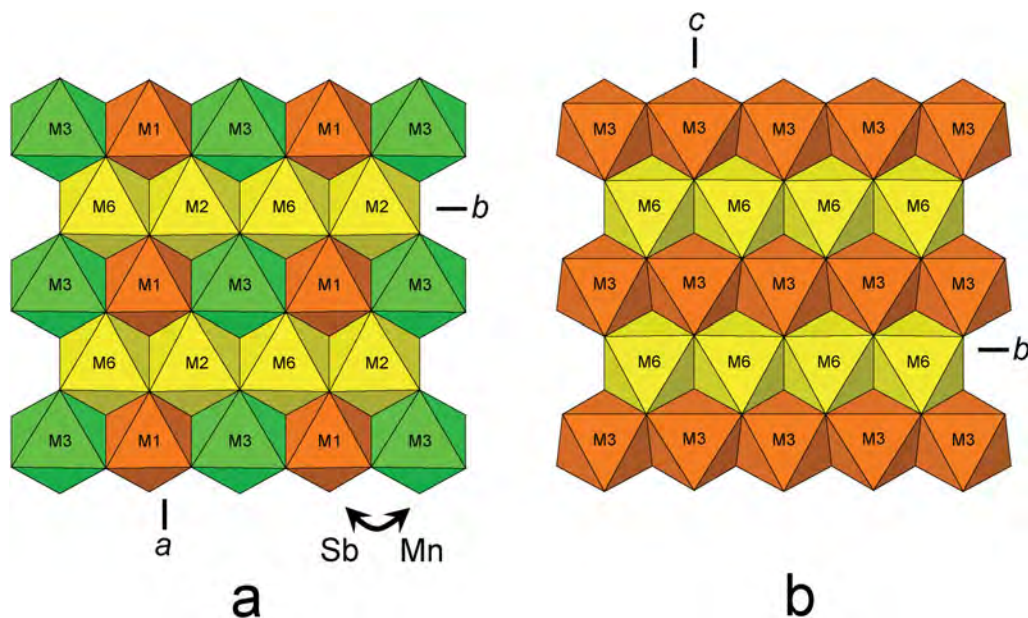


FIG. 5. The F-wall in: (a) folvikite and (b) Sb-rich pinakiolite. Yellow =  $(Mg > Mn^{2+})$ ; green =  $(Mn^{2+} > Mg)$ ; and orange =  $Sb^{5+}$  dominant.

same for both folvikite (18.15 wt.%; Table 1) and Sb-rich pinakiolite (18.09 wt.%; Hansen *et al.*, 1988).

### C-walls

The C-walls for folvikite and Sb-rich pinakiolite are shown in Figs. 6a and 6b, with simplified idealized  $(\text{Mn}^{2+}, \text{Mg})$  octahedra [ $X(1), X(4)$  in folvikite;  $M(4), M(5)$  in Sb-rich pinakiolite] occupying the disordered region between the  $\text{Mn}^{3+}$ -dominant strips of octahedra [ $M(4), M(5)$  in folvikite;  $M(1), M(2)$  in Sb-rich pinakiolite]. Although there are differences between the individual sites of the two structures (e.g. details of Jahn-Teller polyhedron distortions,  $\text{Mn}^{3+}$  content,  $\text{Mn}^{2+}/\text{Mg}$  ratio, occupancy of positionally disordered sites), the basic chemical prevalence at analogous sites is similar: the Jahn-Teller-distorted octahedra [ $M(4), M(5)$  in folvikite;  $M(1), M(2)$  in Sb-rich pinakiolite] are dominated by  $\text{Mn}^{3+}$ , and the intervening disordered cations [ $X(1)$  to  $X(5)$  in folvikite;  $M(4), M(5), M(45)$  in Sb-rich pinakiolite] are dominantly ( $\text{Mn}^{2+} > \text{Mg}$ ), although folvikite is significantly enriched in  $\text{Mn}^{2+}$  relative to Sb-rich pinakiolite.

### S-columns

The S-columns in both folvikite [ $M(7)$  and  $M(8)$ ] and Sb-rich pinakiolite [ $M(7)$ ] are dominated by Mg octahedra; however, folvikite contains  $\sim 2/3$  Mg +  $\sim 1/3$   $\text{Mn}^{2+}$  at these octahedra, compared to octahedra fully occupied by Mg in Sb-rich pinakiolite.

### Connectivity

The connectivity between F-walls, C-walls, S-columns and  $\text{BO}_3$  groups is shown in Fig. 7 for folvikite (Fig. 7a) and Sb-rich pinakiolite (Fig. 7b). The general topology is the same for both minerals. The unit-cell translation ‘up-the-page’ for folvikite is  $1/2$  that of Sb-rich pinakiolite. In Sb-rich pinakiolite, successive C-walls along  $[100]$  are displaced by  $1/2y$  relative to each other, such that  $\text{Mn}^{3+}$  octahedral strips alternate with disordered  $\text{Mn}^{2+}$ –Mg strips in a given (010) section (Fig. 7b). In folvikite, successive C-walls along  $[001]$  are not offset from each other in the (010) plane. Note that the  $M(1)\text{Sb}^{5+}\text{--}M(3)\text{Mn}^{2+}$  ordering within the F-wall of folvikite is not apparent in Fig. 7a because only  $M(1)\text{Sb}^{5+}$  octahedra are present at the (010) level shown (cf. Fig. 5a).

## The chemical formula of pinakiolite, Sb-rich pinakiolite and folvikite

### Pinakiolite

The IMA Master List [<http://nrmima.nrm.se/imalist.htm>], gives the formula for pinakiolite as  $(\text{Mg}, \text{Mn})_2(\text{Mn}^{3+}, \text{Sb}^{5+})\text{O}_2\text{BO}_3$ , which is conveniently collapsed to a compact  $\text{O}_2(\text{BO}_3)$  anion component ( $Z=8$ ). However, this formula lacks utility for Sb-bearing pinakiolites, as the  $\text{Sb}^{5+}$  content must be zero in order to maintain charge balance. The pinakiolite structure contains seven cation sites (other than B), of which three  $Mn(1\text{--}3)$  are clearly occupied by dominant  $\text{Mn}^{3+}$  and four  $Mg(1\text{--}4)$  by dominant Mg, providing  $\text{Mn}^{3+}\text{Mg}_2$  pfu (Moore and Araki, 1974). The empirical formula for pinakiolite given by Moore and Araki (1974) is:



The pinakiolite formula can be written as  $\text{Mn}^{3+}(\text{Mg}, \text{Mn}^{2+})_2\text{O}_2(\text{BO}_3)$  ( $Z=8$ ) for Sb-free pinakiolite.

### Sb-rich pinakiolite

In Sb-rich pinakiolite, the  $M(3)$  site occupancy is reported as  $(\text{Sb}_{0.38}\text{Mn}_{0.32}\text{Mg}_{0.30})$  with Sb dominant at the site. However, it was stated that the Mn and Mg contents at  $M(3)$  were not determined reliably (Norrestam and Hansen, 1990). The  $M(3)$  site has a  $\langle M(3)\text{--O} \rangle$  distance of 2.05 Å, and applying our general relation between bond length and constituent radii  $\langle M\text{--O} \rangle = 0.834(r) + 1.497$  from folvikite, gives a constituent radius of 0.66 Å for the  $M(3)$  site in Sb-rich pinakiolite, which can be regarded as a good approximation for this homologous structure. Thus, the relatively short  $\langle M(3)\text{--O} \rangle$  distance of 2.05 Å is not compatible with significant occupancy by the larger  $\text{Mn}^{2+}$  cation (0.83 Å radius, Shannon, 1976), and any Mn content must be dominated by  $\text{Mn}^{3+}$  (0.645 Å radius), as in pinakiolite. The individual  $M(3)\text{--O}$  bond lengths show very little variation in Sb-rich pinakiolite, and as  $\text{Mn}^{3+}$  is expected to induce bond-length distortions associated with Jahn-Teller effects, the absence of such a distortion indicates little or no  $\text{Mn}^{3+}$  at the  $M(3)$  site. The site occupancy reported for  $M(3)$  by Norrestam and Hansen (1990) is  $(\text{Sb}_{0.38}\text{Mn}_{0.32}\text{Mg}_{0.30})$  which corresponds to a site-scattering of 30.98 electrons per site. If we exclude Mn from the  $M(3)$  site, this site-scattering value is concordant with the site occupancy  $(\text{Mg}_{0.51}\text{Sb}_{0.49})$  which has an aggregate

# FOLVIKITE, A NEW OXYBORATE MINERAL

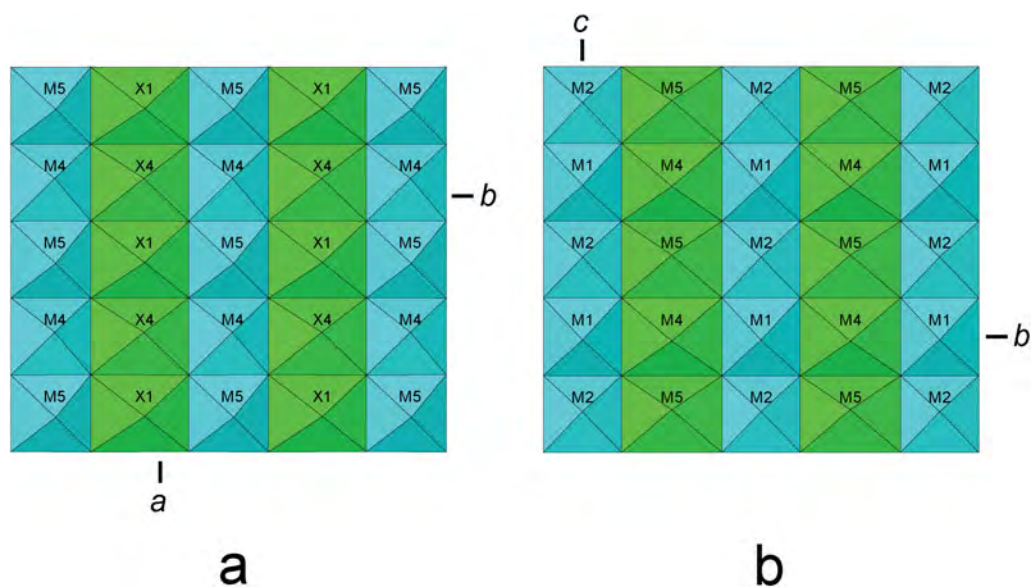


FIG. 6. The C-walls in: (a) folvikite and (b) Sb-rich pinakiolite. Blue =  $\text{Mn}^{3+}$  dominant and green = ( $\text{Mn}^{2+} > \text{Mg}$ ).

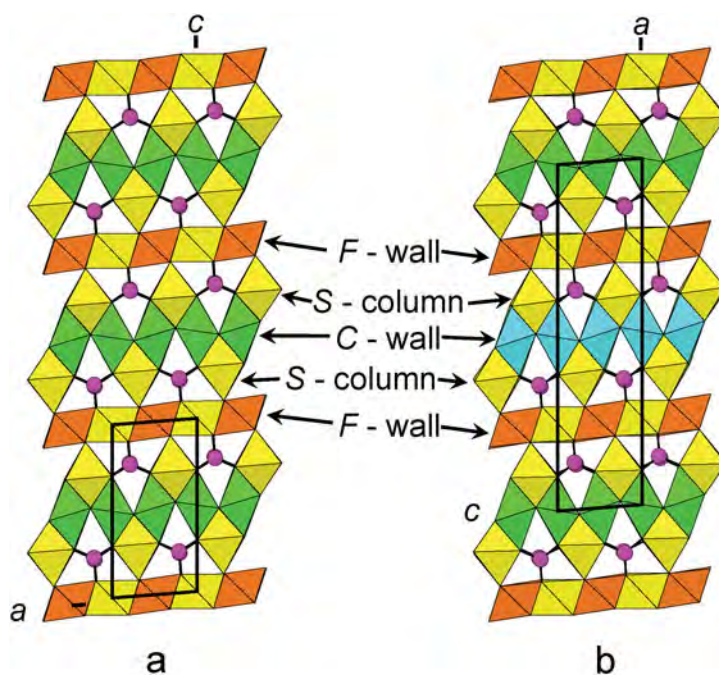


FIG. 7. The crystal structure of: (a) folvikite projected down  $[010]$ ; (b) Sb-rich pinakiolite projected down  $[010]$ . Legend as in Figs 5, 6; pink circles = B atoms.

TABLE 8. The differences between the *M* sites in Sb-rich pinakiolite and pinakiolite.

<i>M</i> sites	<i>M</i> (3)	<i>M</i> (1,2)	<i>M</i> (6,7)	<i>M</i> (4,45,5)	Anions	<i>Z</i>
Sb-rich pinakiolite	Mg,Sb <sup>5+</sup>	Mn <sup>3+</sup>	Mg <sub>3</sub>	Mn <sup>2+</sup> ,Mg	O <sub>4</sub> (BO <sub>3</sub> ) <sub>2</sub>	4
Pinakiolite	Mn <sub>2</sub> <sup>3+</sup>			Mg <sub>4</sub>	O <sub>4</sub> (BO <sub>3</sub> ) <sub>2</sub>	4

ionic radius of 0.66 Å. We therefore suggest that only Sb<sup>5+</sup> and Mg are present at *M*(3), and that all Mn<sup>3+</sup> fully occupies the *M*(1) and *M*(2) sites, resulting in Mn<sub>0.5</sub><sup>3+</sup> pfu (*Z* = 8). The Sb-rich pinakiolite formula calculated using the electron-microprobe results of Hansen *et al.* (1988) with *B* = 1, Mn<sup>3+</sup> = 0.5 and 5 anions pfu (*Z* = 8) gives:



The Sb<sub>0.238</sub><sup>5+</sup> content (pfu; *Z* = 8), if entirely ordered at the *M*(3) site, would result in an Sb<sup>5+</sup> site-occupancy of 0.476, which is in good agreement

with the (Mg<sub>0.51</sub> Sb<sub>0.49</sub>) site occupancy proposed above. Sb-rich pinakiolite is derived from pinakiolite via Sb<sup>5+</sup> + Mg replacing Mn<sup>3+</sup> at the *M*(3) site, with additional homovalent Mn<sup>2+</sup> → Mg substitution occurring at the disordered *M* sites within the C-wall. An *M*(3) site fully occupied by Mn<sup>3+</sup> (i.e. pinakiolite) would require substitution by 1/3 Sb<sup>5+</sup> + 2/3 Mg to maintain charge balance, but the proposed *M*(3) site content of ~ Sb<sub>0.5</sub> exceeds this value, and an additional heterovalent substitution at another site(s) is required to achieve charge balance; this might involve divalent cations (Mn<sup>2+</sup>, Mg) at the trivalent (Mn) *M*(1) and *M*(2) sites, and/

TABLE 9. Comparison of selected properties for folvikite and Sb-rich pinakiolite.

	Folvikite	Sb-rich pinakiolite
Locality	Nordmark, Sweden	Långban, Sweden
Habit	striated prisms	tabular platy crystals
Colour	black to reddish brown	light olive-green to yellow-brown
Lustre	submetallic	semimetallic
Space group	<i>P</i> 2	<i>C</i> 2/ <i>m</i>
<i>a</i> (Å)	5.3767(10)	21.773(4)
<i>b</i>	6.2108(10)	6.153(2)
<i>c</i>	10.9389(18)	5.327(2)
β (°)	94.399(9)	94.38(2)
<i>V</i> (Å <sup>3</sup> )	364.22(16)	711.6(4)
<i>Z</i>	1	4
Density <sub>calc</sub> (g cm <sup>-3</sup> )	4.14	3.92
Sb <sub>2</sub> O <sub>5</sub>	18.15	18.09
Mn <sub>2</sub> O <sub>3</sub>	11.62	18.54*
Fe <sub>2</sub> O <sub>3</sub>	0.45	0.06
Al <sub>2</sub> O <sub>3</sub>	0.27	0.22
MnO	29.73	14.93*
MgO	24.11	32.02
B <sub>2</sub> O <sub>3</sub> **	15.27	16.36
Total	99.60	100.22
Formula	Sb <sup>5+</sup> Mn <sup>3+</sup> (Mg,Mn <sup>2+</sup> ) <sub>10</sub> O <sub>8</sub> (BO <sub>3</sub> ) <sub>4</sub>	(Mg,Sb <sup>5+</sup> )Mn <sup>3+</sup> Mg <sub>3</sub> (Mn <sup>2+</sup> ,Mg)O <sub>4</sub> (BO <sub>3</sub> ) <sub>2</sub>
References	This work	Hansen <i>et al.</i> (1988) Norrestam and Hansen (1990)

\* Mn<sup>3+</sup> / (Mn<sup>2+</sup> + Mn<sup>3+</sup>) adjusted to provide Mn<sup>3+</sup> = 1 apfu; \*\* calculated using *B* = 4 (folvikite) and *B* = 2 apfu (Sb-rich pinakiolite).



or vacancies at the disordered  $M(4,45,5)$  sites within the C-wall. If we constrain the  $\text{Sb}^{5+}$  to occur at  $M(3)$  in subordinate fashion following the exchange  $\frac{1}{3} \text{Sb}^{5+} + \frac{2}{3} \text{Mg} \rightarrow \text{Mn}^{3+}$ , we can convey the difference between pinakiolite and Sb-rich pinakiolite in the formula unit by reducing the  $Z$  value to 4, as shown in Table 8. The F-wall contains the  $M(3)$  and  $M(6)$  sites, the C-wall contains the ordered  $M(1)$  and  $M(2)$  sites and the disordered  $M(4-5)$  sites, and the S-column contains the  $M(7)$  site.

### Folvikite

The  $P2$  structure of folvikite contains a distinct Sb-dominant site, and the empirical formula  $(\text{Sb}_{1.02}^{5+}\text{Mn}_{1.34}^{3+}\text{Mg}_{5.46}\text{Mn}_{2.82}^{2+}\text{Al}_{0.05}\text{Fe}_{0.26}^{3+})_{\Sigma 12}\text{O}_8(\text{BO}_3)_4$  can be simplified to  $\text{Sb}^{5+}\text{Mn}^{3+}(\text{Mg}, \text{Mn}^{2+})_{10}\text{O}_8(\text{BO}_3)_4$  ( $Z = 1$ ) to show the  $\text{Sb}^{5+}$  site-dominant character. In comparing folvikite with Sb-rich pinakiolite: (1) both have similar  $\text{Sb}_2\text{O}_5$  content; (2) folvikite contains less  $\text{Mn}_2\text{O}_3$  and  $\text{MgO}$ , but more  $\text{MnO}$ , relative to Sb-rich pinakiolite.

As is apparent from the above discussion, folvikite and Sb-rich pinakiolite are structurally and chemically quite similar; Table 9 compares selected properties of these two minerals. However, their habit and colour are quite distinct. Folvikite occurs as striated prisms whereas Sb-rich pinakiolite occurs as tabular platy crystals. Moreover, folvikite is black to reddish brown whereas Sb-rich pinakiolite is light olive-green to yellow-brown; presumably the additional Mg in Sb-rich pinakiolite serves to inhibit charge-transfer effects between heterovalent transition-metals.

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### Supplementary material

To view supplementary material for this article, please visit <https://doi.org/10.1180/minmag.2017.081.059>

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